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Mantle and crust interaction in post-collisional setting,  
a case study of lamprophyric-granitic composite dykes  
of northern Aigoual pluton, French Massif Central

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# Summary

- **Introduction**
- Objectives and Methodology
- Results and Discussion

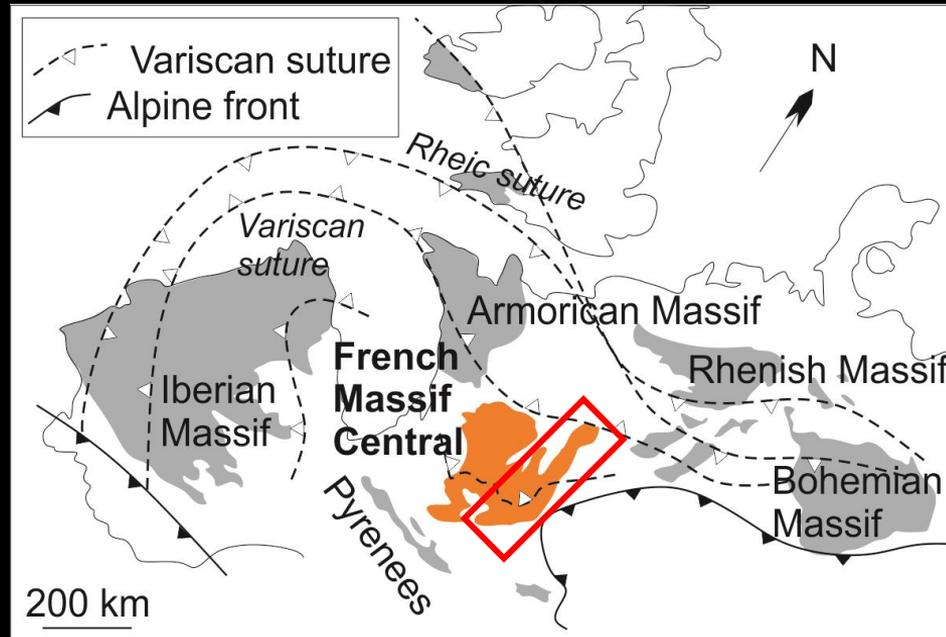
Chapter II – paper in preparation to submission

Chapter III – general discussion

- Conclusions

# Introduction

- Variscan belt – Gondwana + Laurussia = Pangea
- French Massif Central – inner part of the Variscan belt, exposing allochthonous and parautochthonous units at different crustal levels



Modified from  
Moyen et al. (2017)

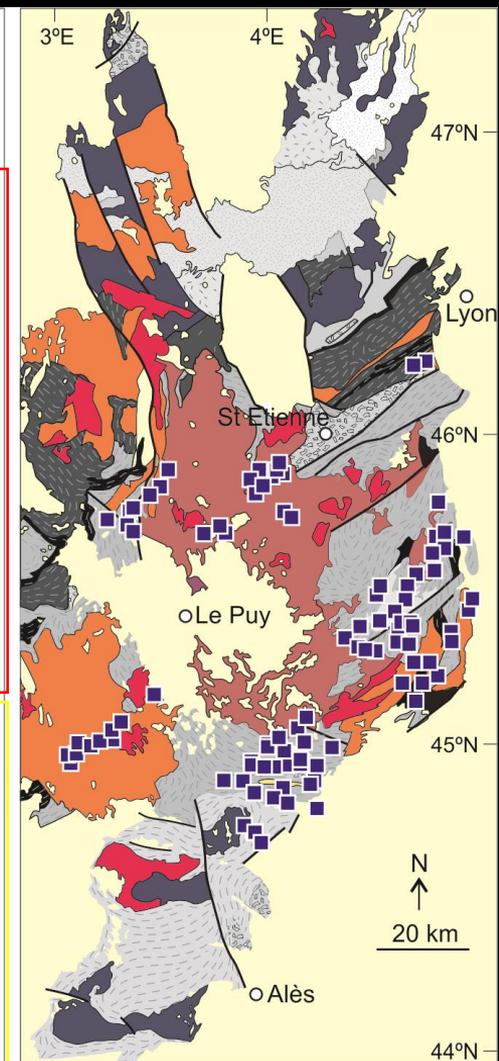
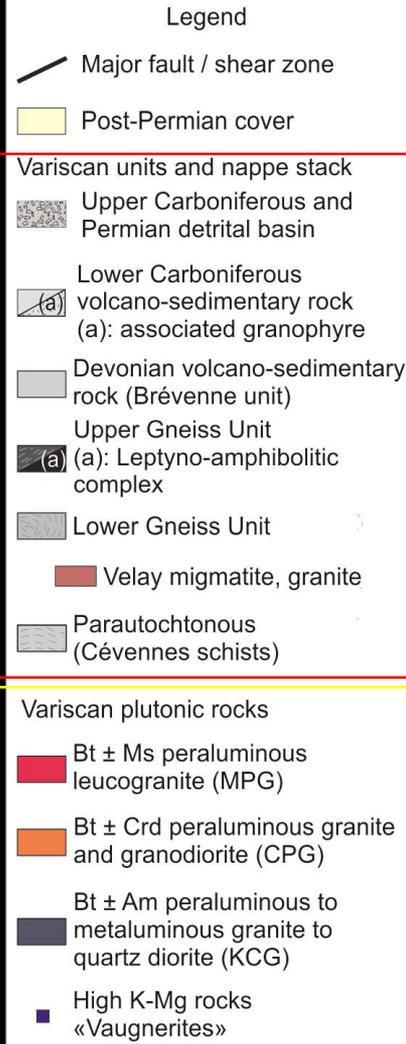
# Ordovician to Carboniferous

## South-verging nappe stack metamorphic units

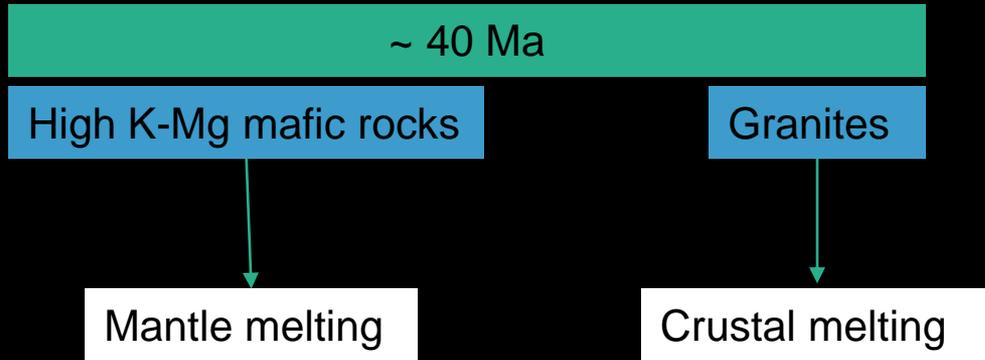
- Opening of oceanic basins (500 – 450 Ma)
- Closure and subduction (420 – 360 Ma)
- Proper collision (350 – 320 Ma)

## Post-collisional magmatic rocks

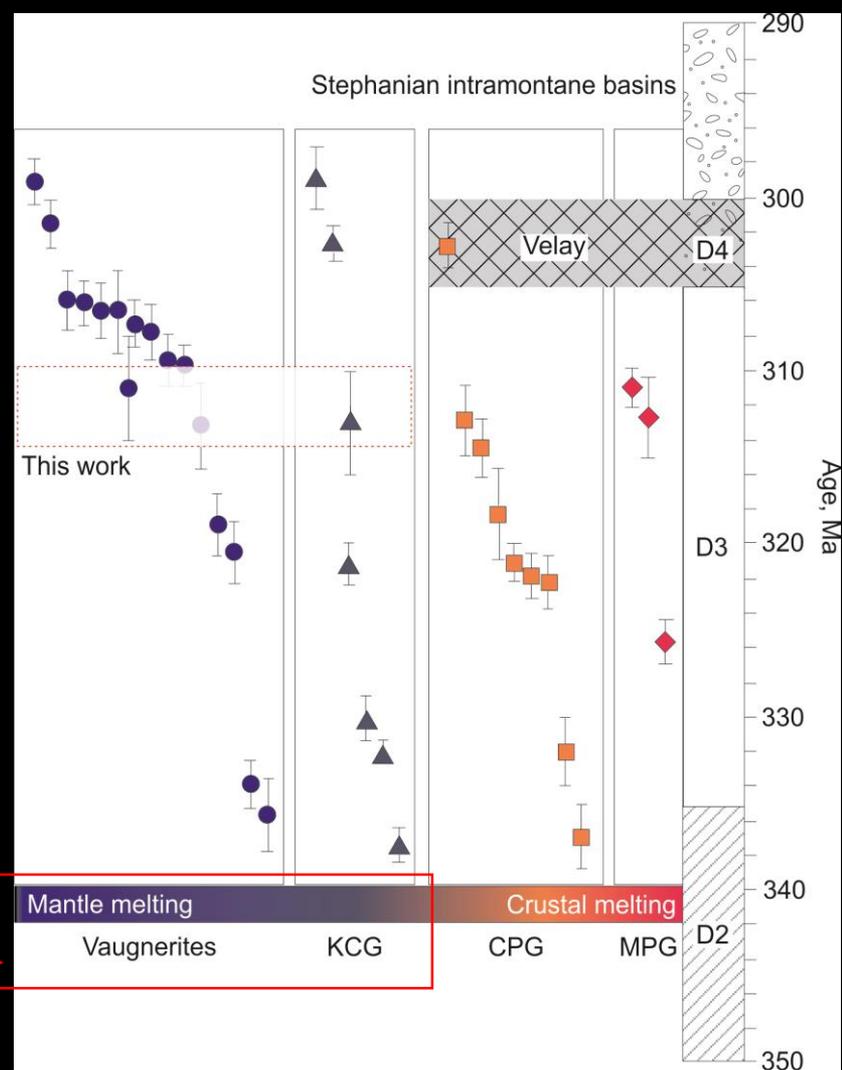
Modified from Couzinié et al. (2016)

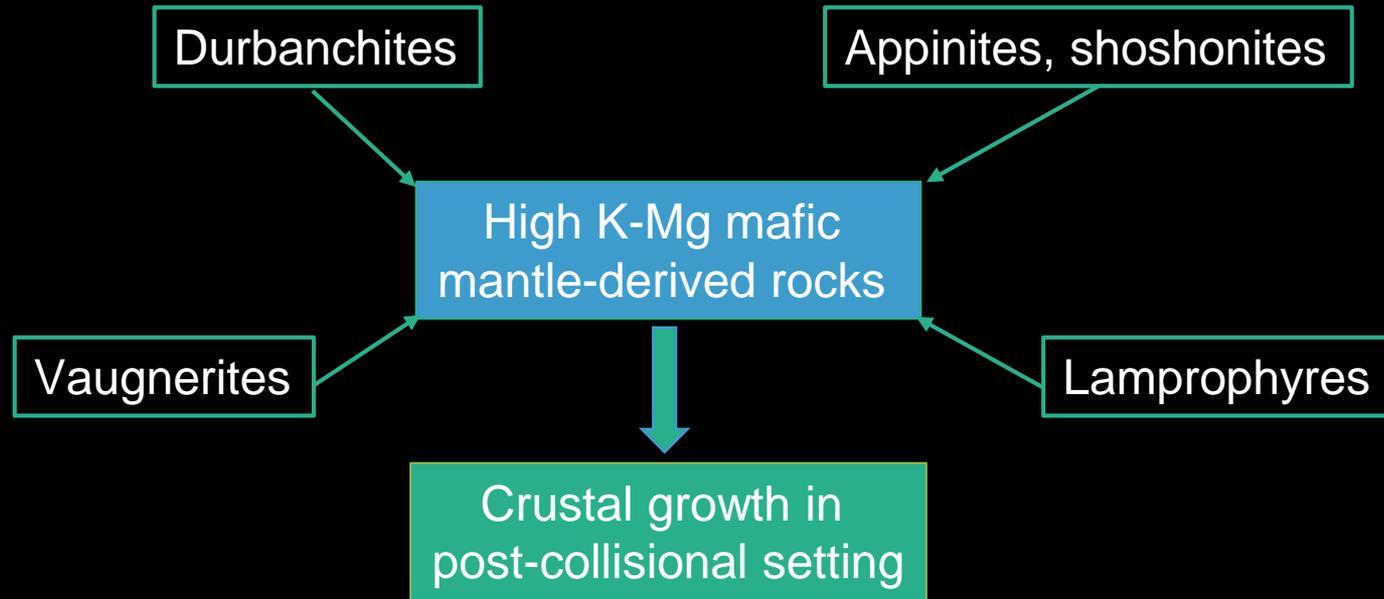


- Extensive and long-lived ~40 Ma magmatism



Modified from Laurent et al. (2017)





Potassic to ultrapotassic diorites, gabbros and/or lamprophyres; metaluminous composition with intermediate  $\text{SiO}_2$  contents (55-70 wt%). (von Raumer et al. 2014; Laurent et al. 2017).

High K-Mg mafic  
mantle-derived rocks

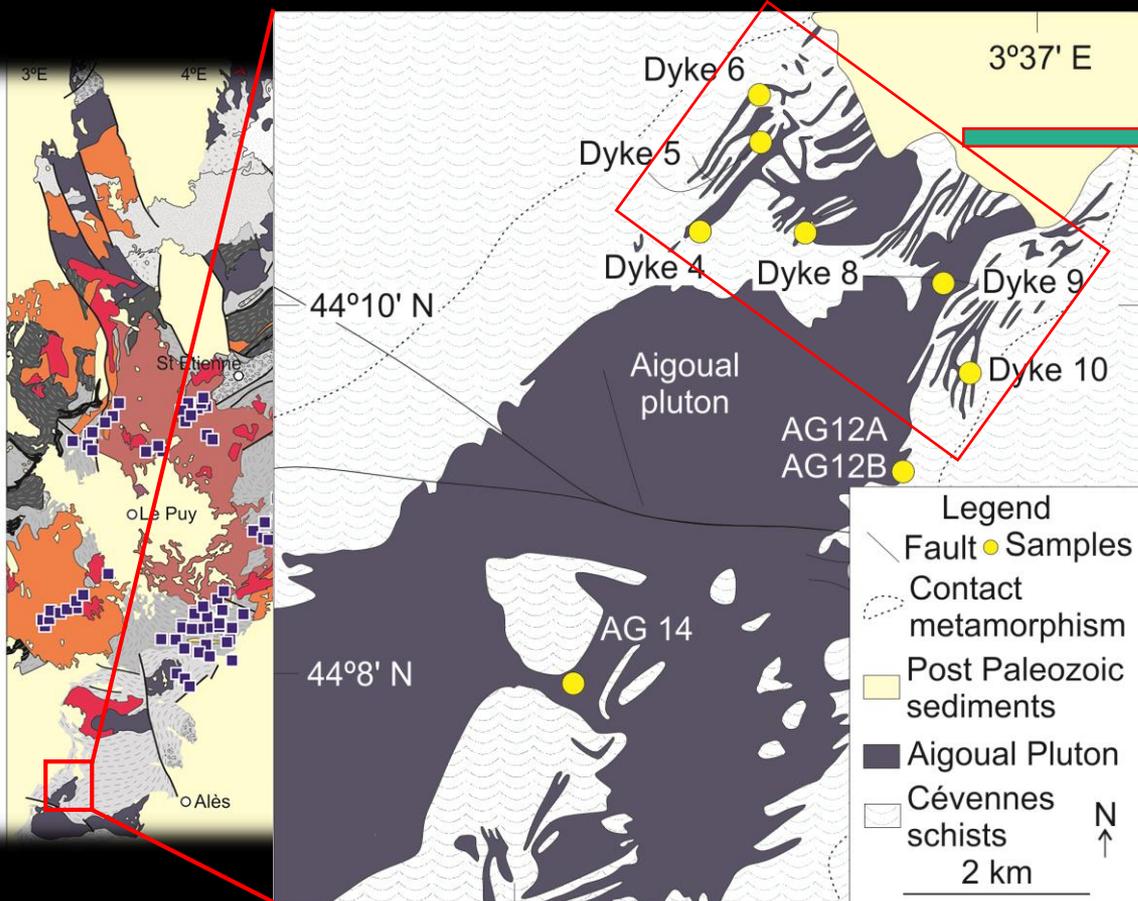


Despite being well defined in space and time, the role of the mafic magmatism in post-Variscan setting is still debated...

- Differentiated magmas from the mafic rocks or the products of hybridization between the latter and crustal melts;
- The balance between crustal growth by addition of new mantle-derived material and recycling of crust in the mantle in post-collisional sites is poorly understood.

# NNE composite dykes on the border of Aigoual pluton

- Legend
- Major fault / shear zone
  - Post-Permian cover
  - Variscan units and nappe stack
    - Upper Carboniferous and Permian detrital basin
    - Lower Carboniferous volcano-sedimentary rock (a): associated granophyre
    - Devonian volcano-sedimentary rock (Brévenne unit)
    - Upper Gneiss Unit
    - (a) Leptyno-amphibolitic complex
    - Lower Gneiss Unit
    - Velay migmatite, granite
    - Parautochthonous (Cévennes schists)
  - Variscan plutonic rocks
    - Bt ± Ms peraluminous leucogranite (MPG)
    - Bt ± Crd peraluminous granite and granodiorite (CPG)
    - Bt ± Am peraluminous to metaluminous granite to quartz diorite (KCG)
    - High K-Mg rocks (Vaugnerites)



Lamprophyre

Granite

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# Objectives and Methodology

- The objective consists in the petrogenesis of the lamprophyres and granites;
- Constraint the mantle contribution in the post-collisional magmatism in FMC and consequently implications for crustal growth.

 Field work and sampling

 Petrography and mineral chemistry

 Whole-rock major and trace elements

 Sr-Nd-Hf isotopes in whole-rock

 U-Pb and Lu-Hf isotopes in zircon

 Oxygen isotopes in quartz and feldspar

# Summary

- Introduction
- Objectives and Methodology
- **Results and Discussion**

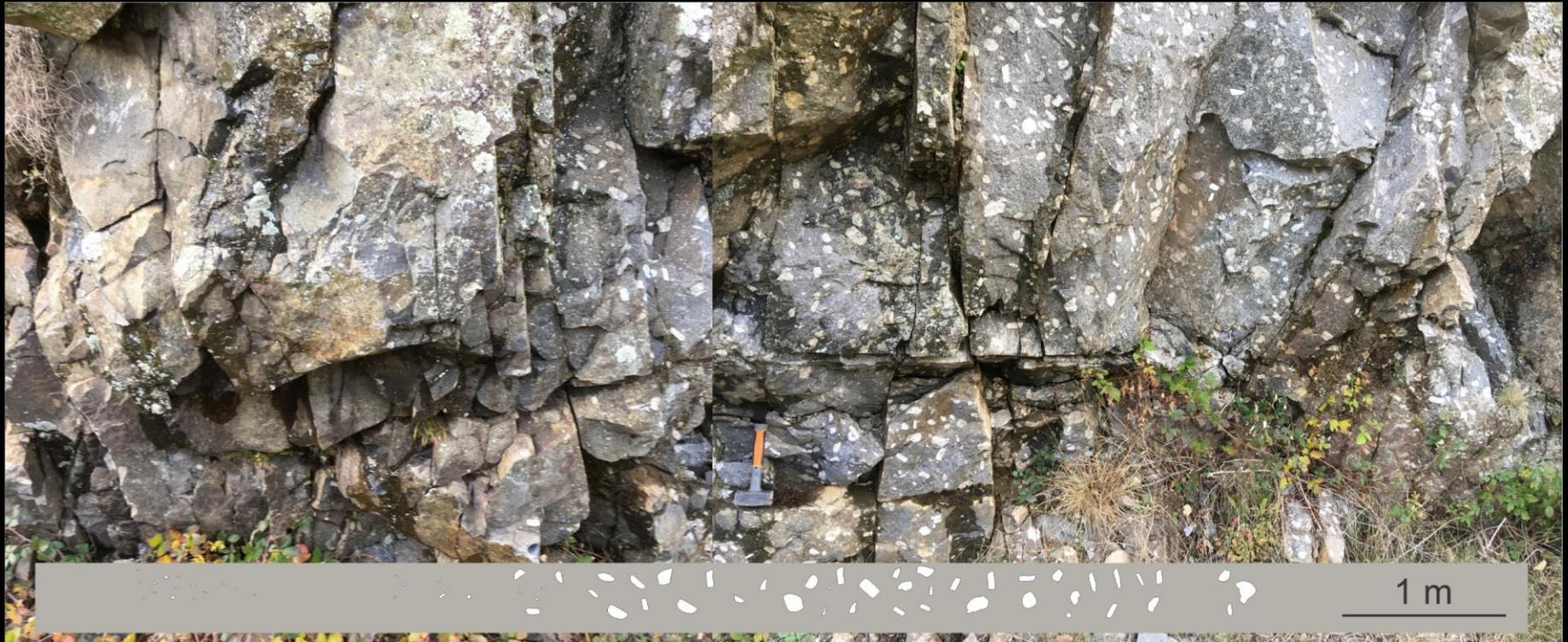
**Chapter II – paper in preparation to submission**

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# Results and Discussion

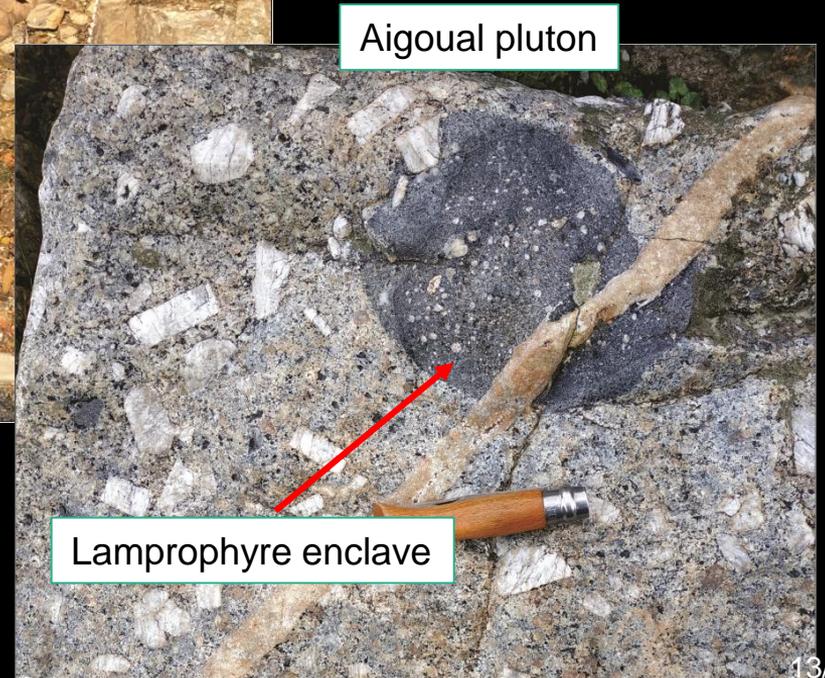
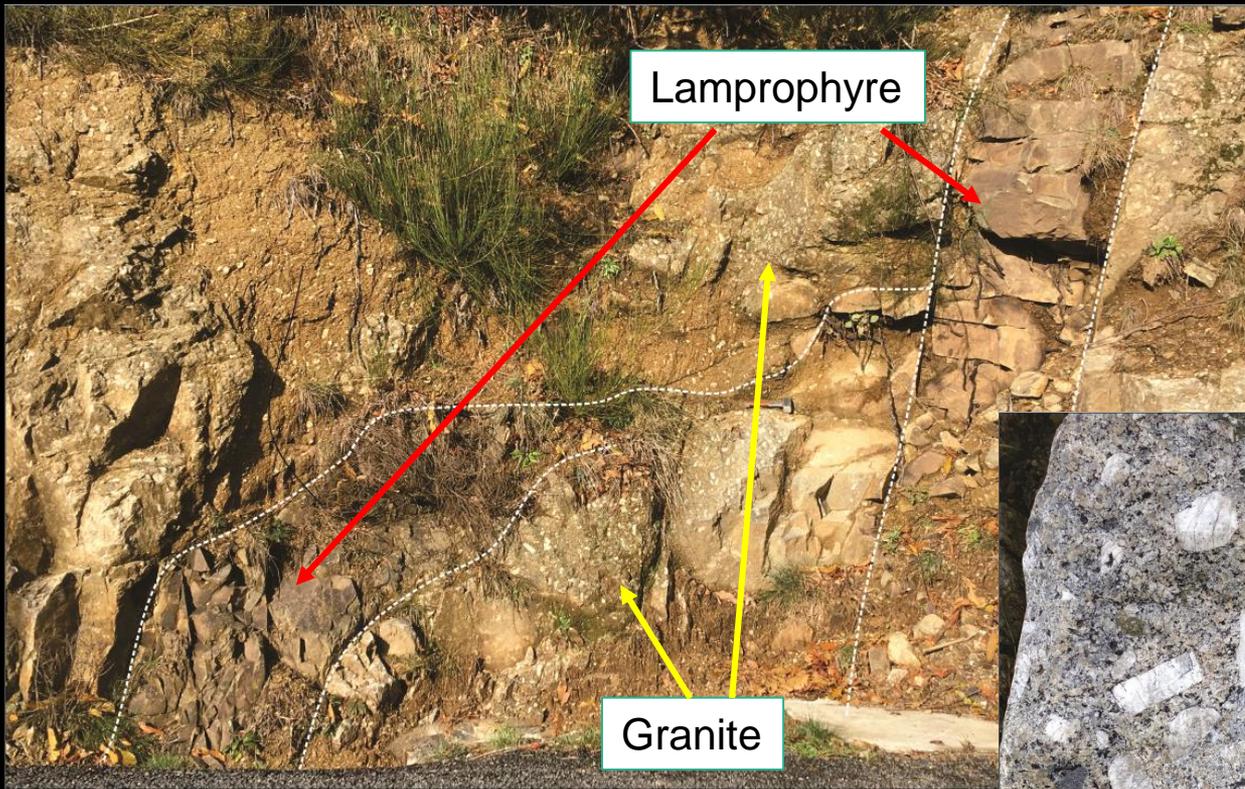
- Field relationships



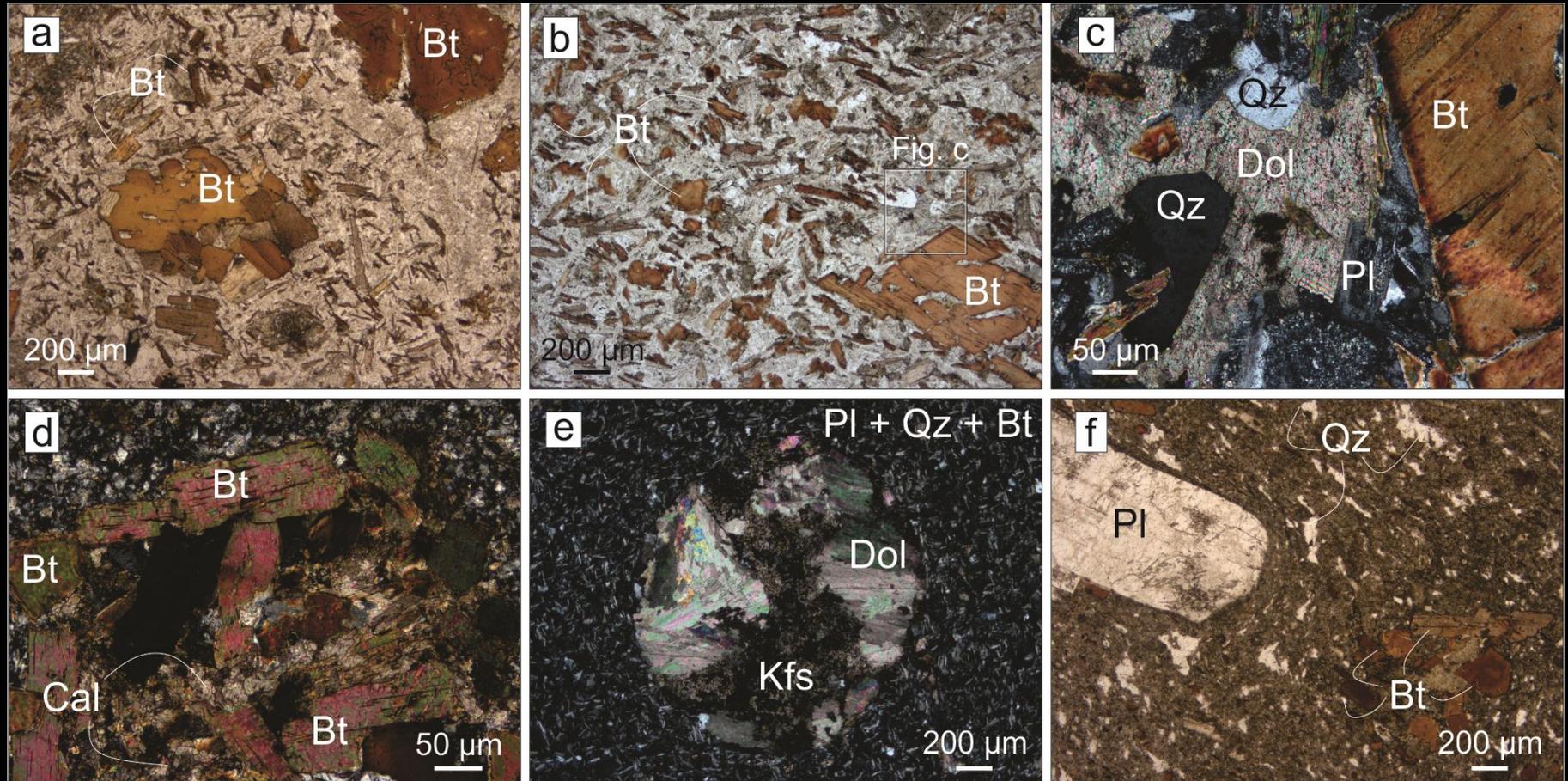
Lamprophyre



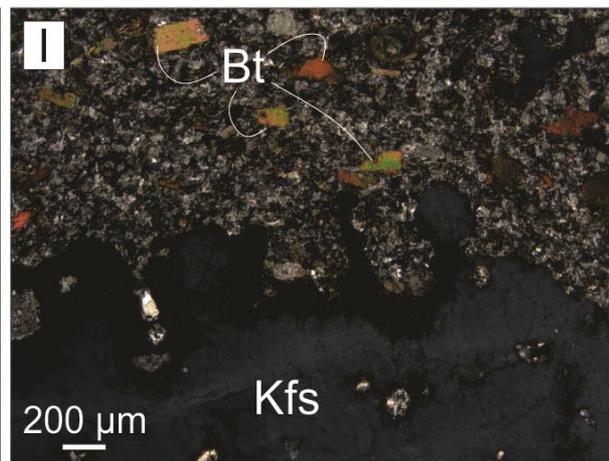
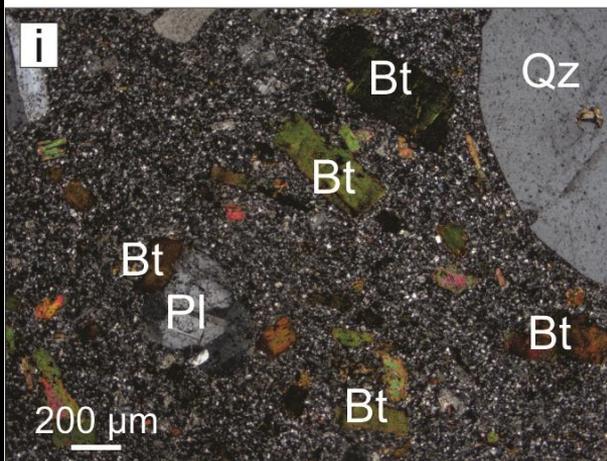
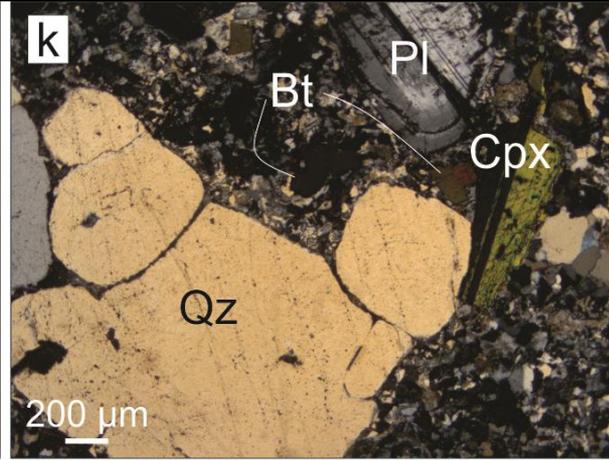
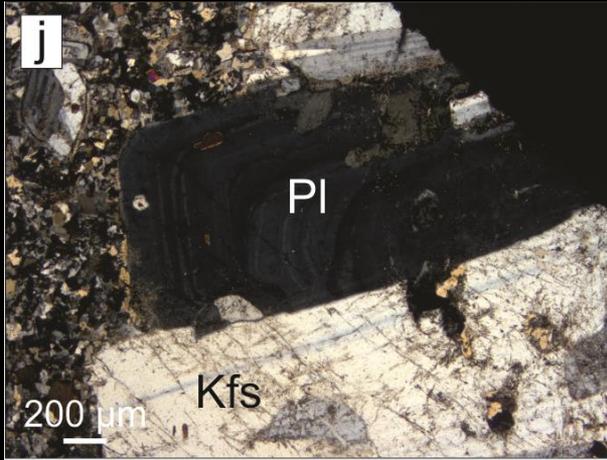
Granite



- Petrography – lamprophyres

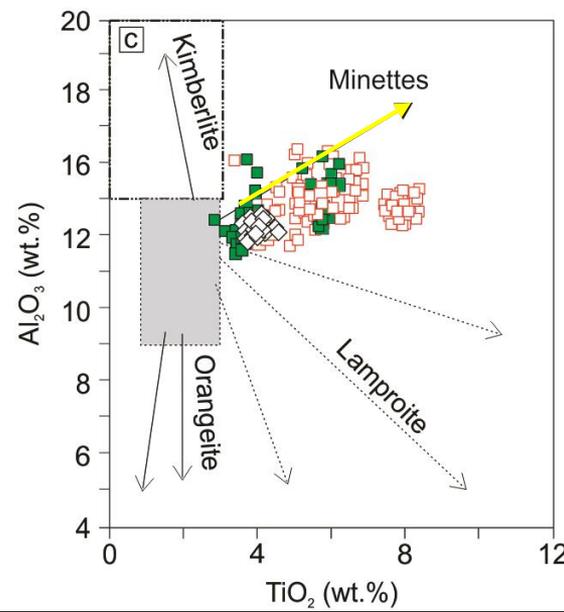
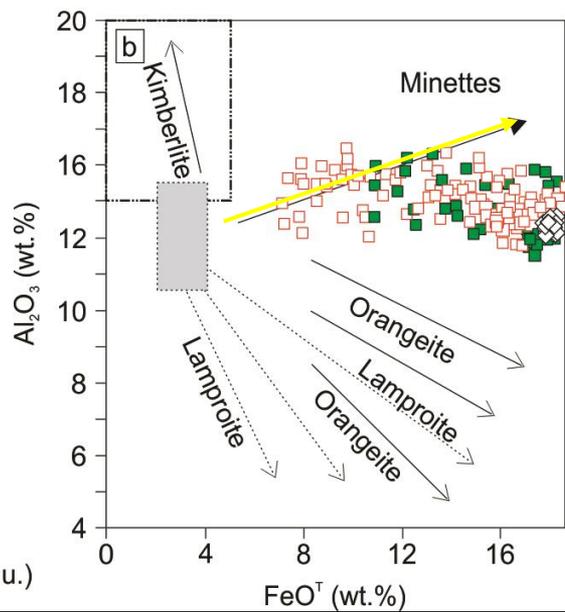
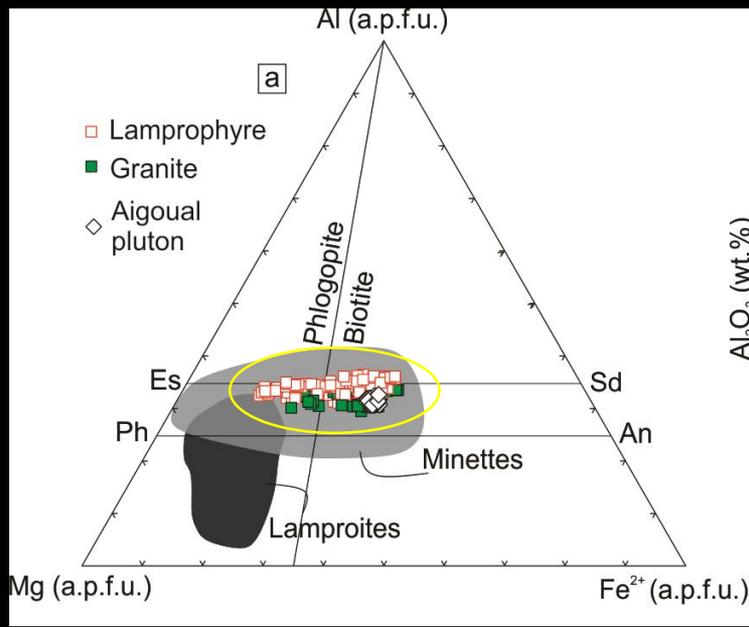
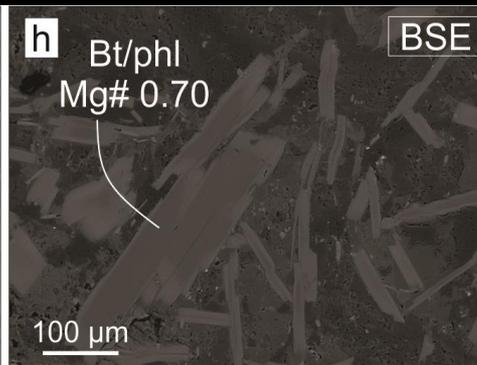
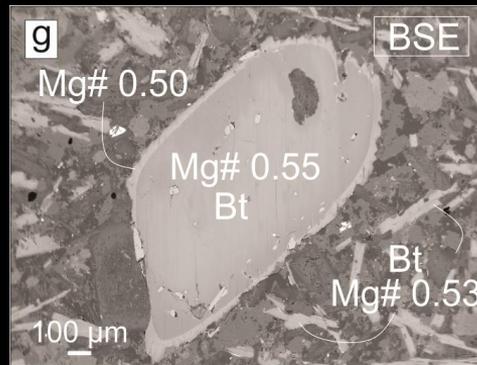


- Petrography – granites

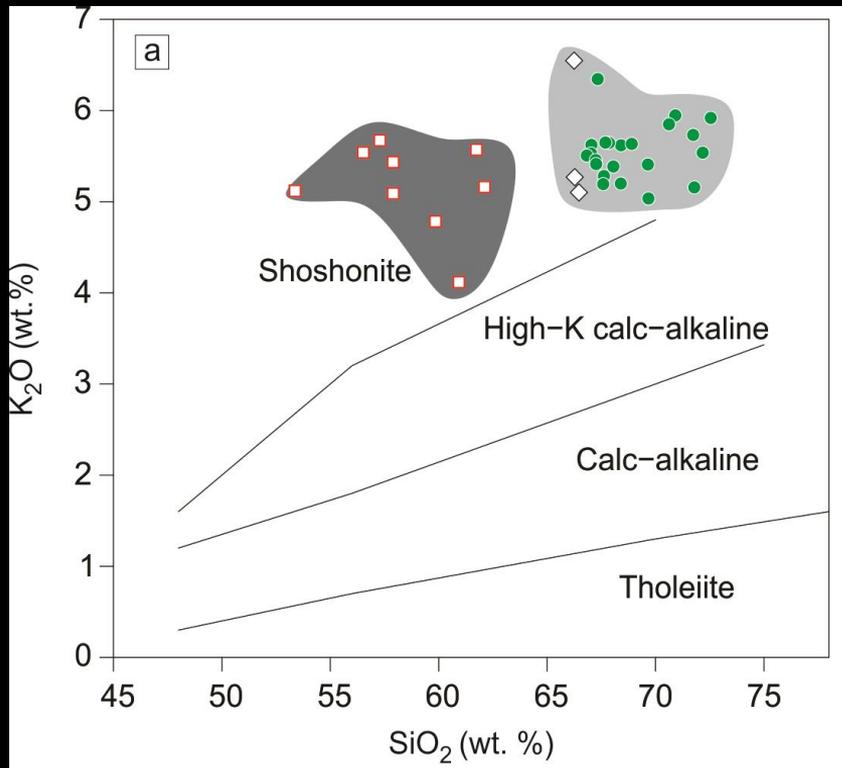


- Mineral chemistry

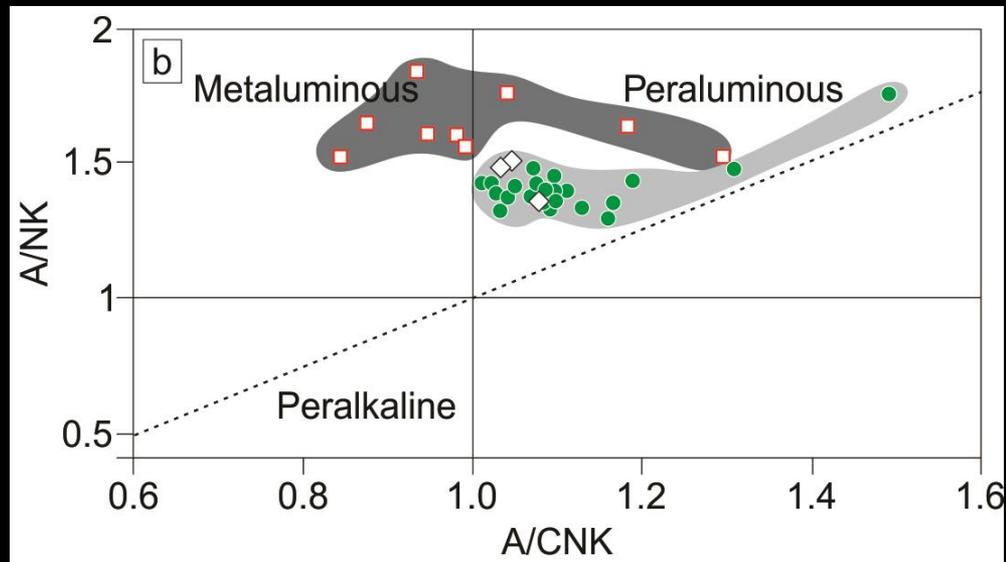
# Mineral chemistry of mica – lamprophyres, granites and Aigoual pluton



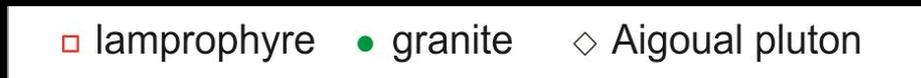
- Whole-rock major elements



From Peccerillo and Taylor (1976)



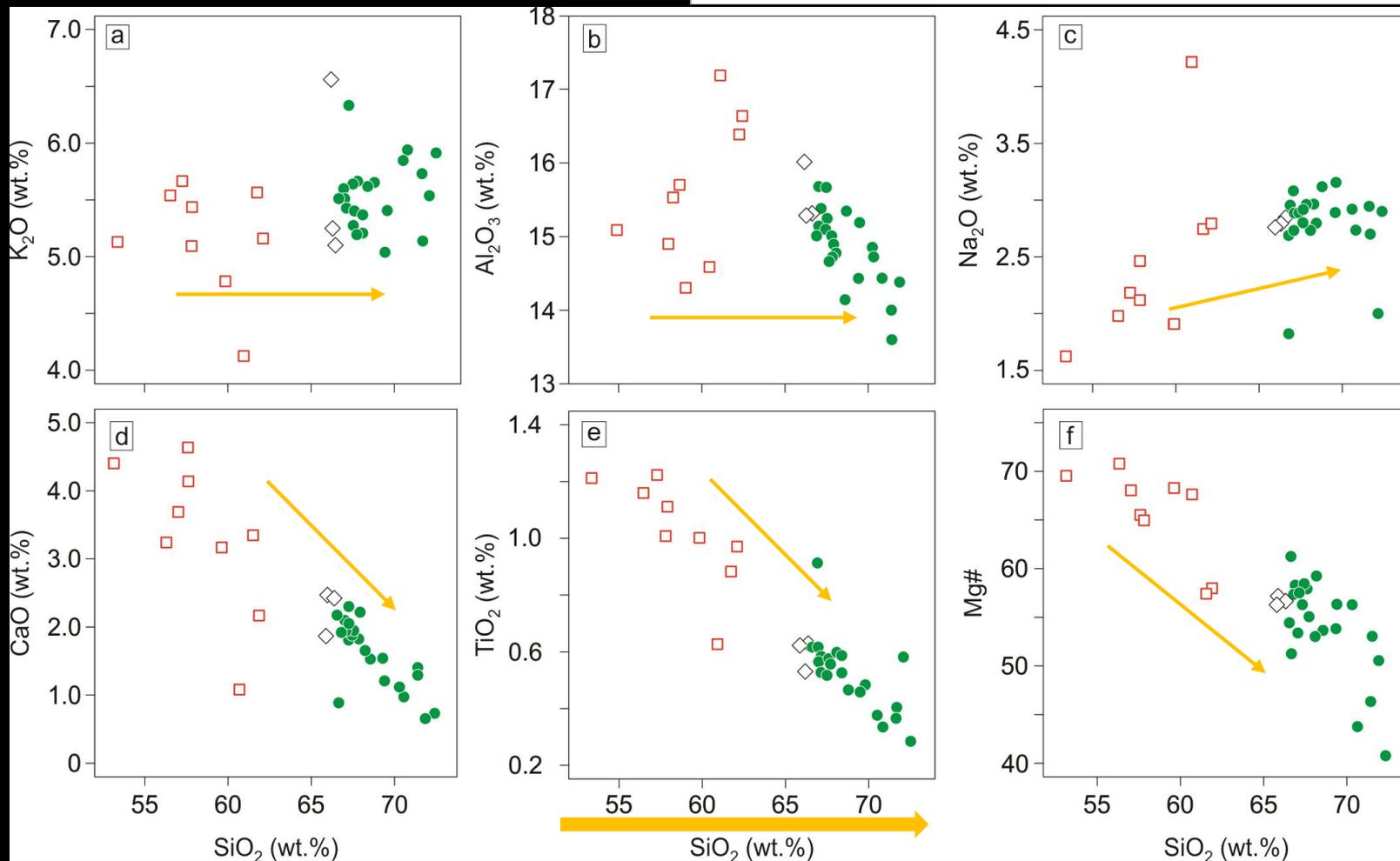
From Shand (1943)

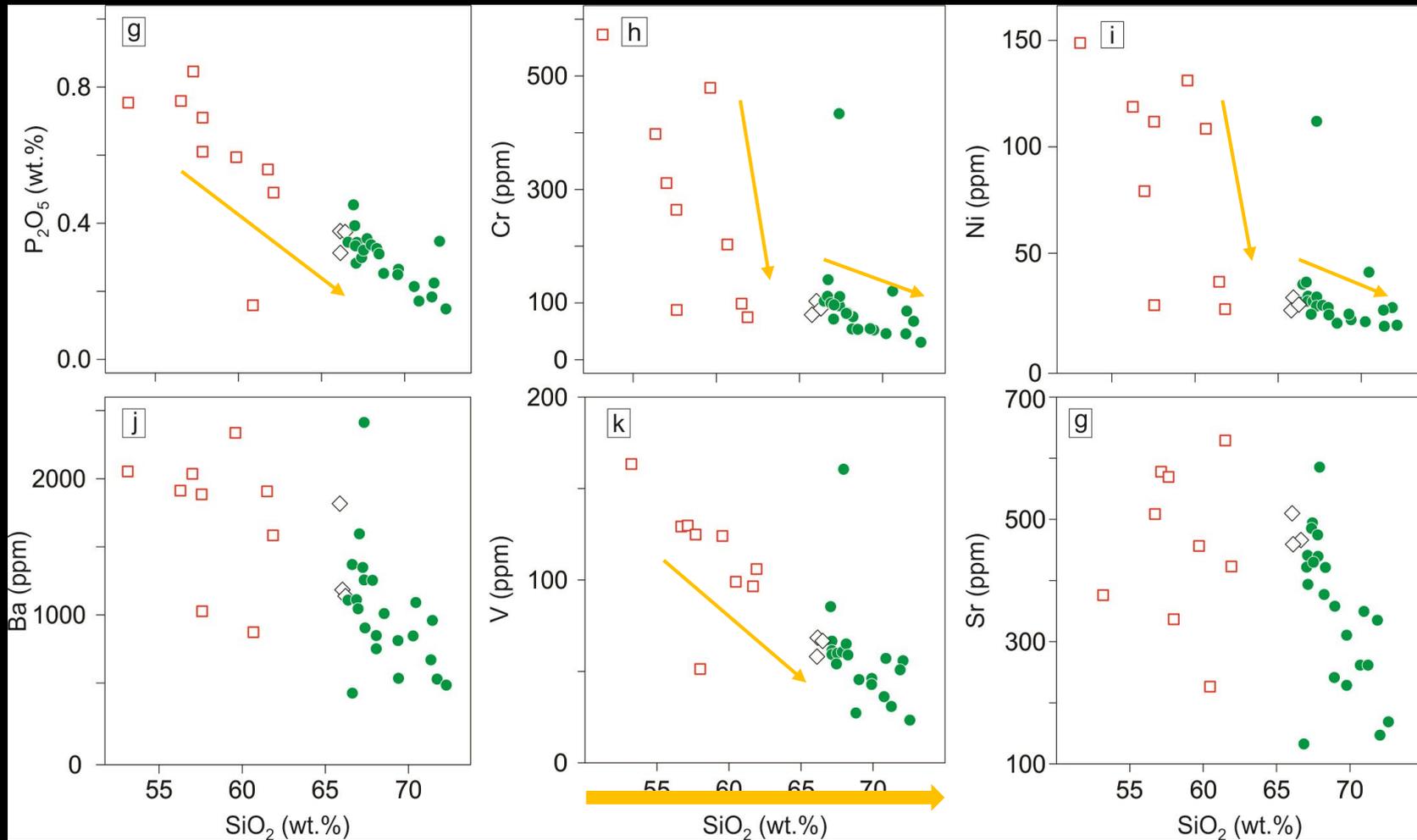


Composite dykes

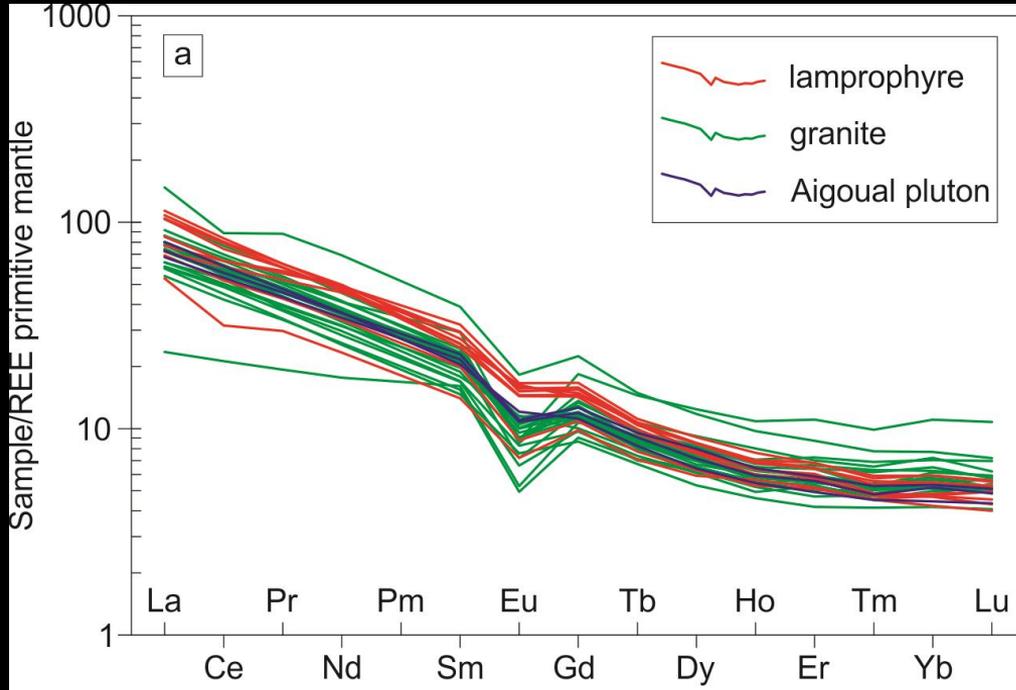
# Harker diagrams

□ lamprophyre   ● granite   ◇ Aigoual pluton

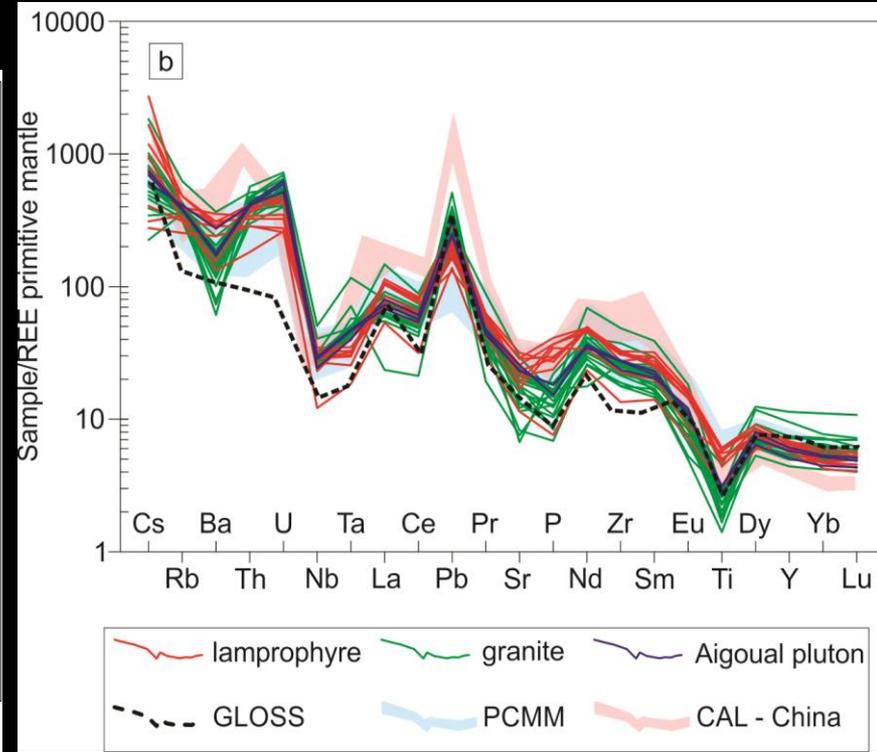




- Trace element patterns



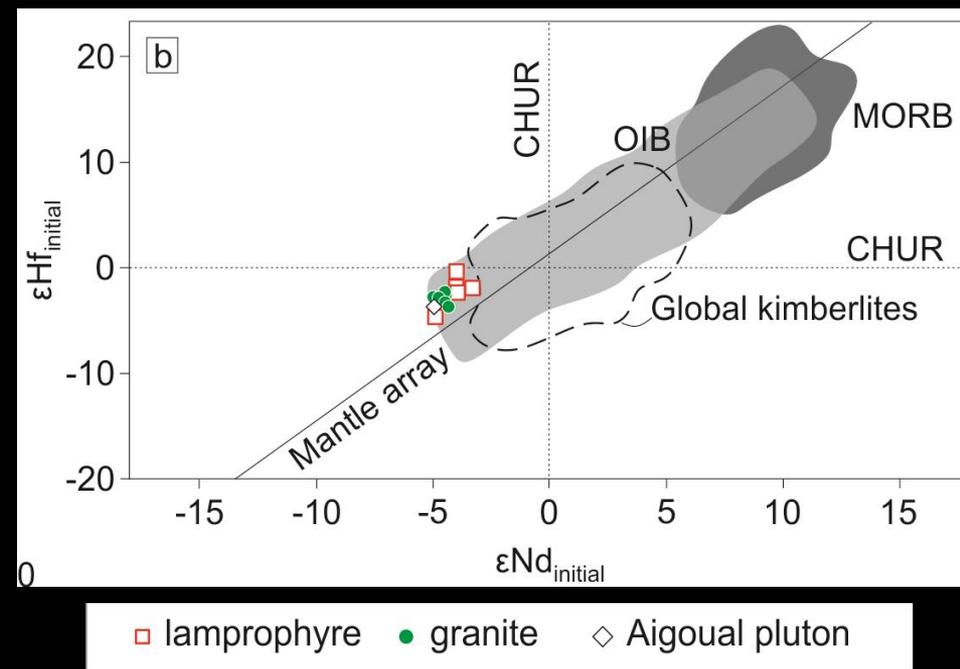
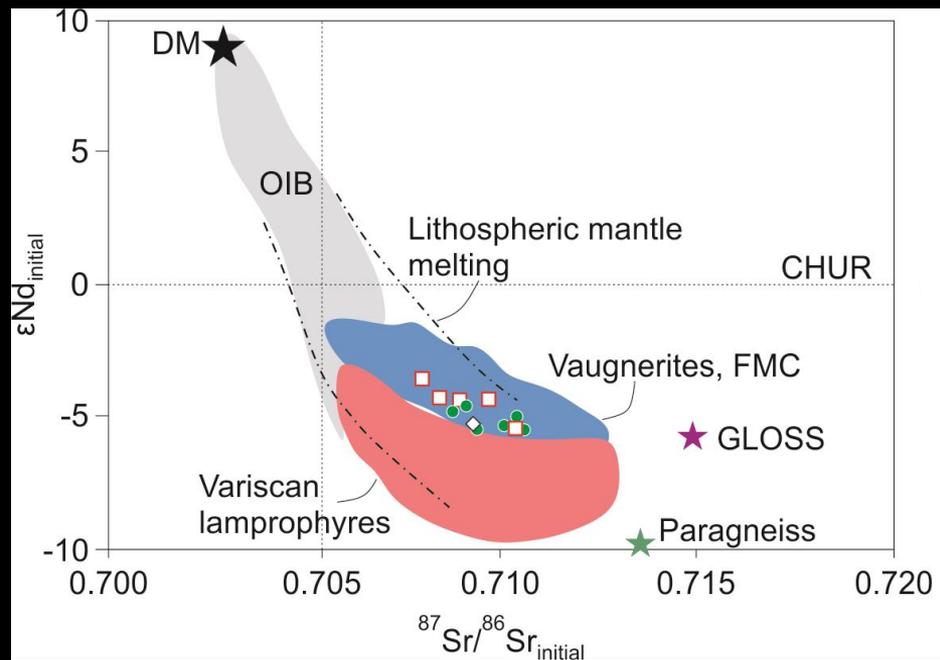
Normalized to the primitive mantle from McDonough and Sun (1995)



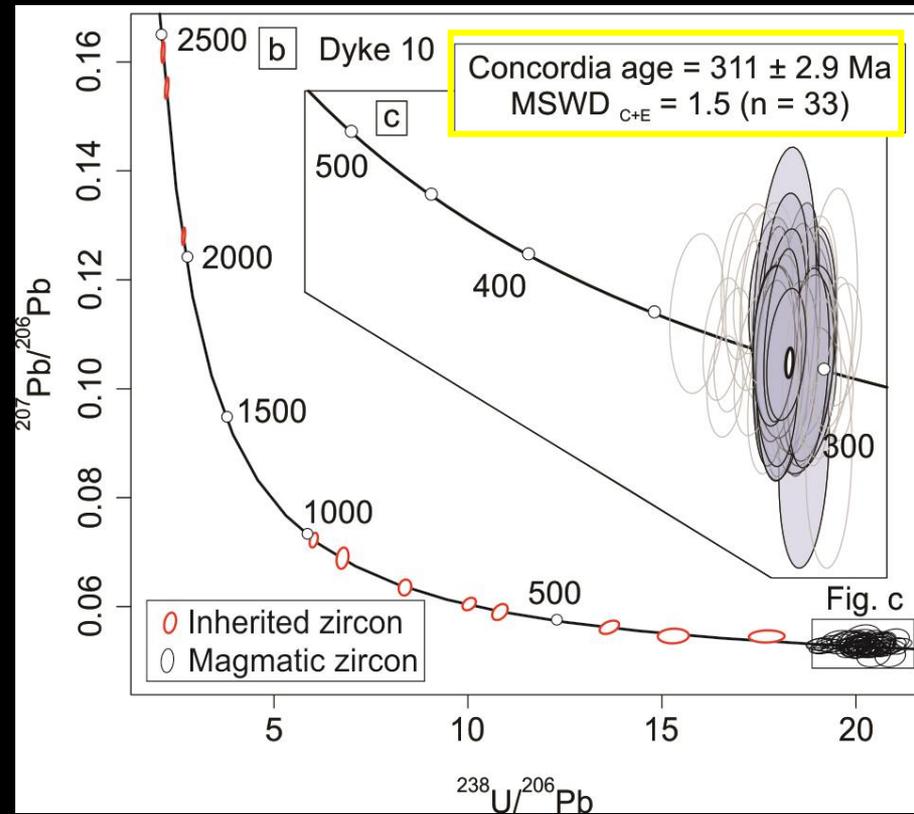
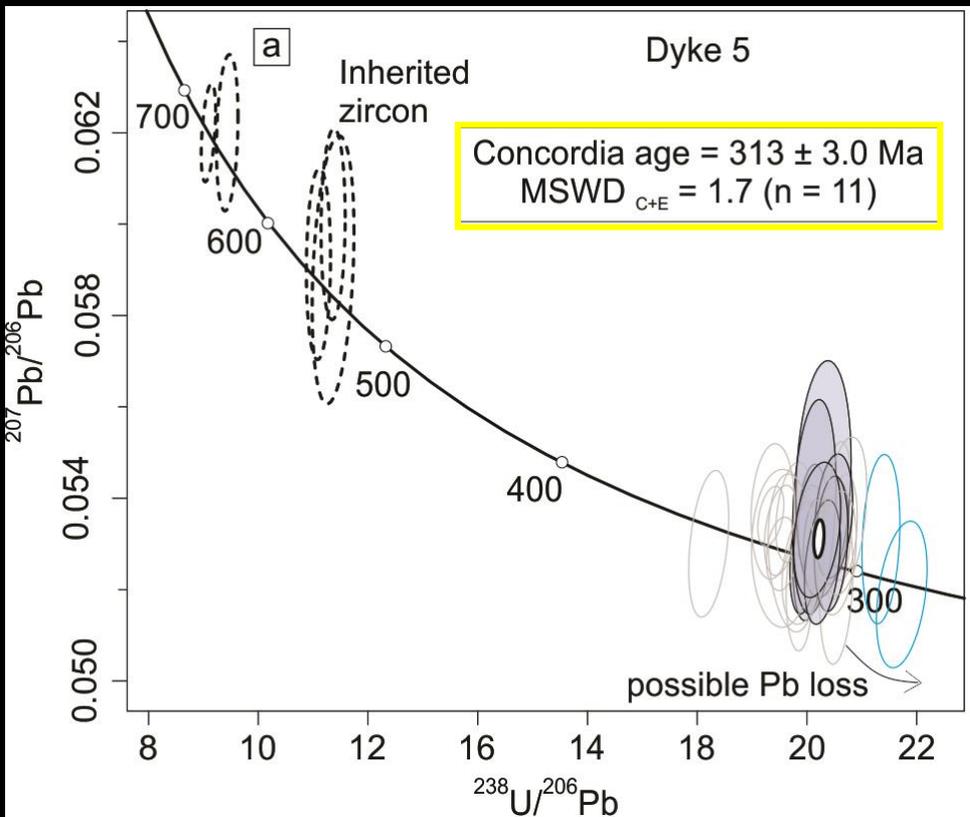
GLOSS - Plank and Langmuir (1998);  
 PCMM – Couzinié et al. (2016);  
 CAL – China – Calc-alkaline lamprophyre from Su et al. (2017).

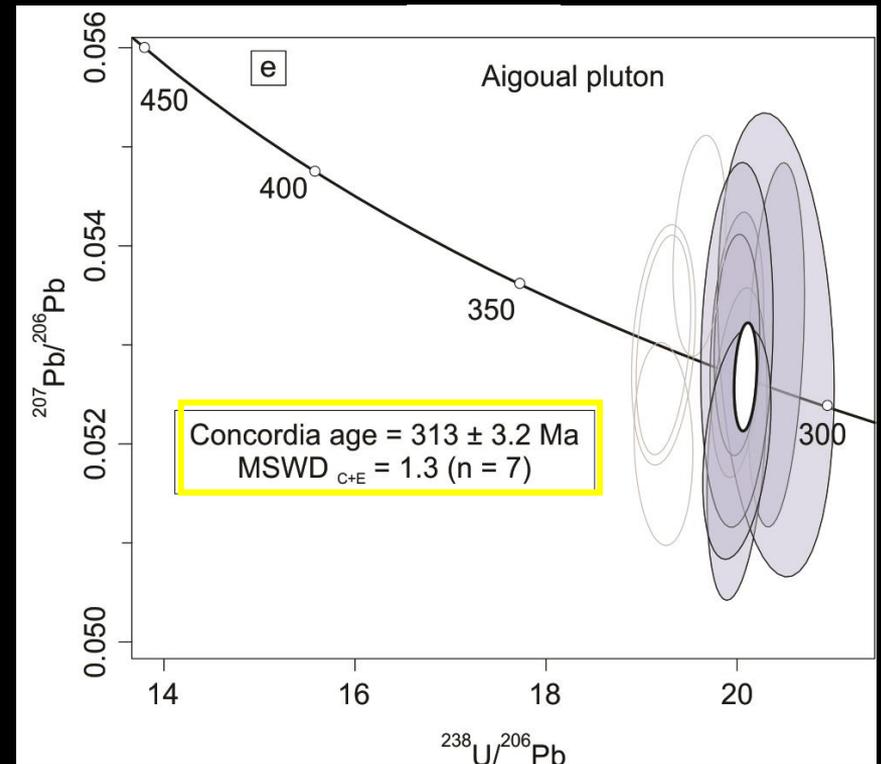
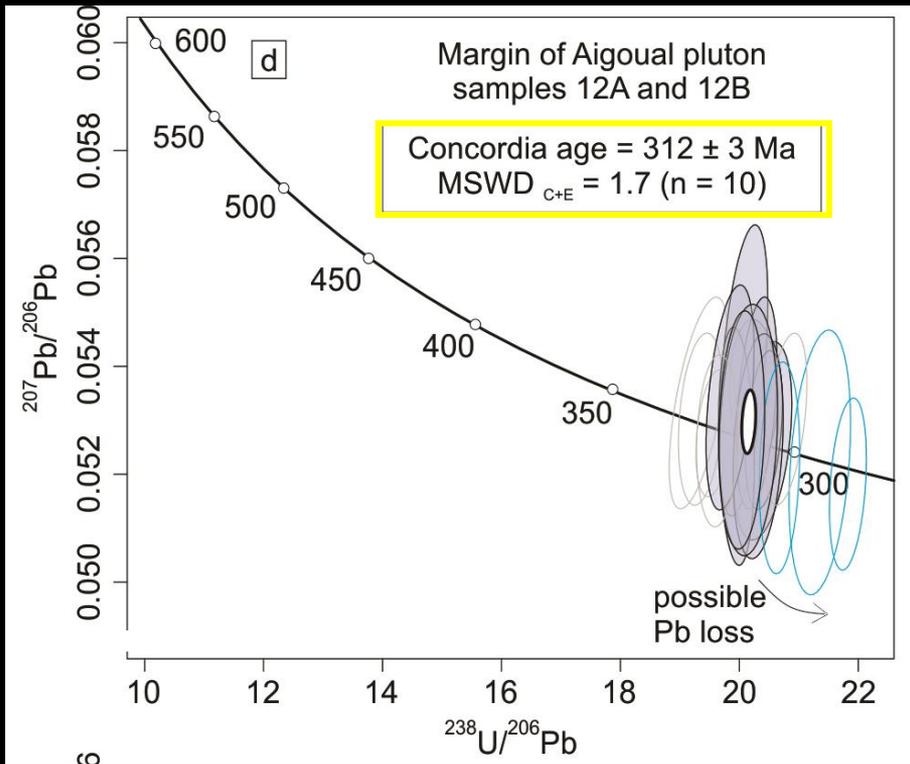
- Whole-rock Sr-Nd-Hf isotopes

- ✓ Lamprophyres and granites display similar  $\epsilon\text{Nd}$  and  $\epsilon\text{Hf}$  and  $^{87}\text{Sr}/^{86}\text{Sr}_{\text{initial}}$
- ✓ Subchondritic  $\epsilon\text{Nd}$  and  $\epsilon\text{Hf}$  and high  $^{87}\text{Sr}/^{86}\text{Sr}_{\text{initial}}$  values



- Zircon U-Pb ages

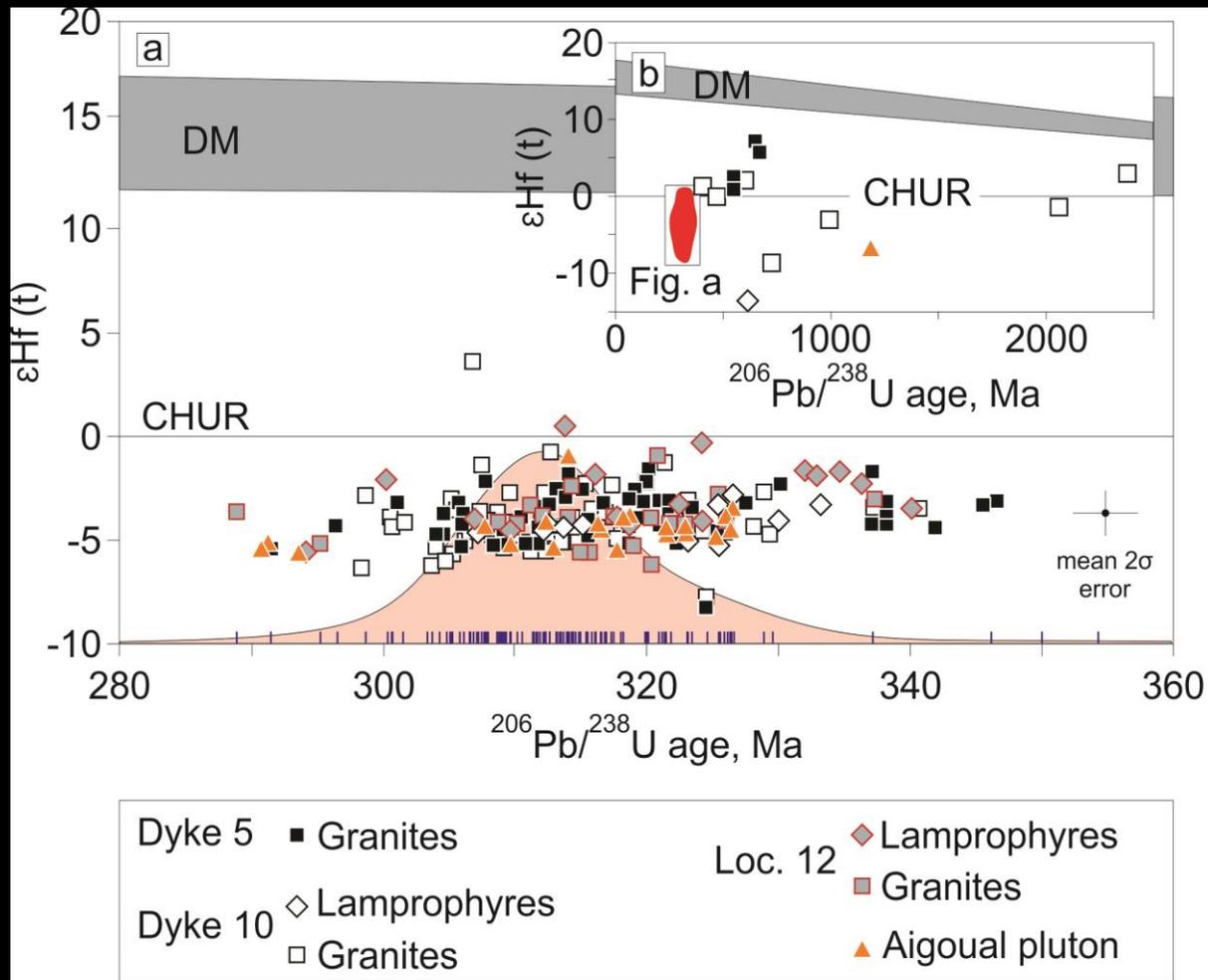




✓ Dykes and the Aigoual pluton show the same crystallization age within uncertainties

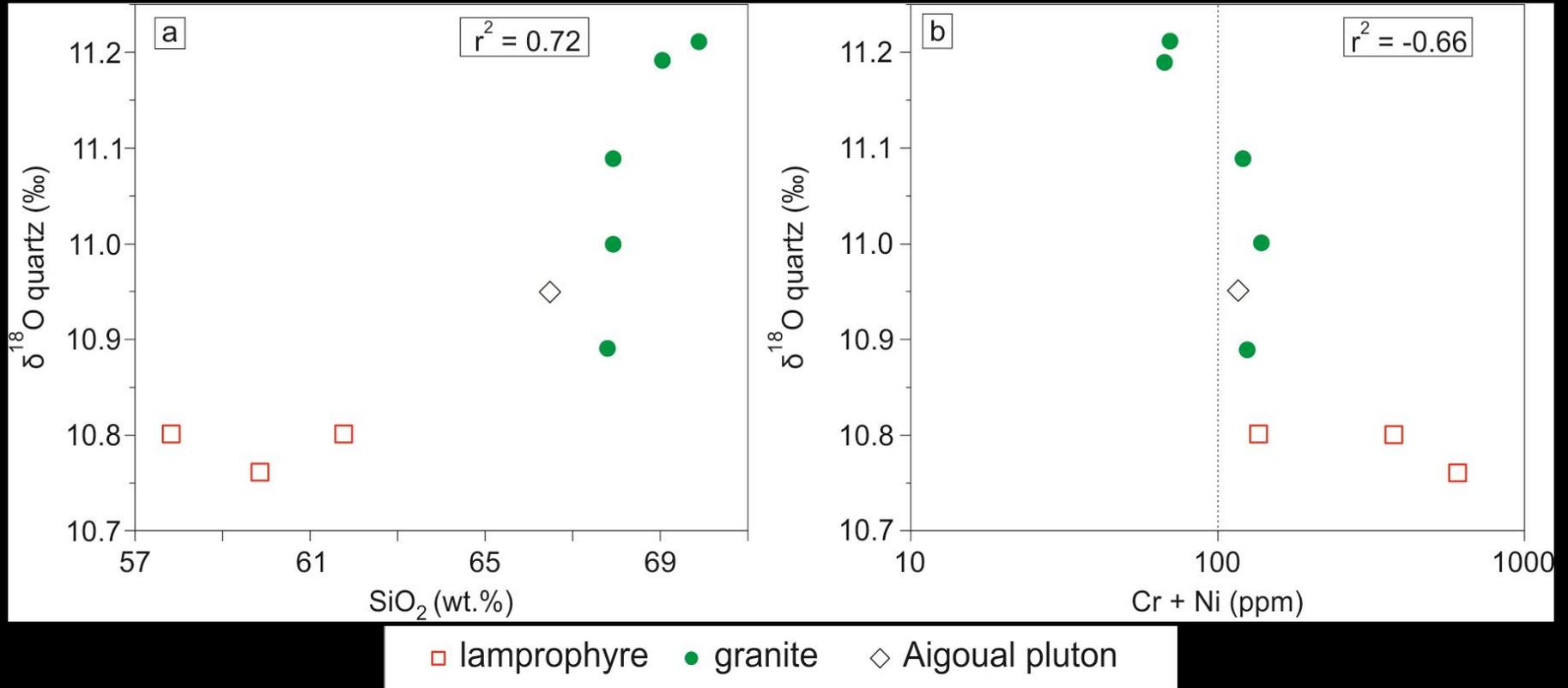
- Zircon Lu-Hf isotopes

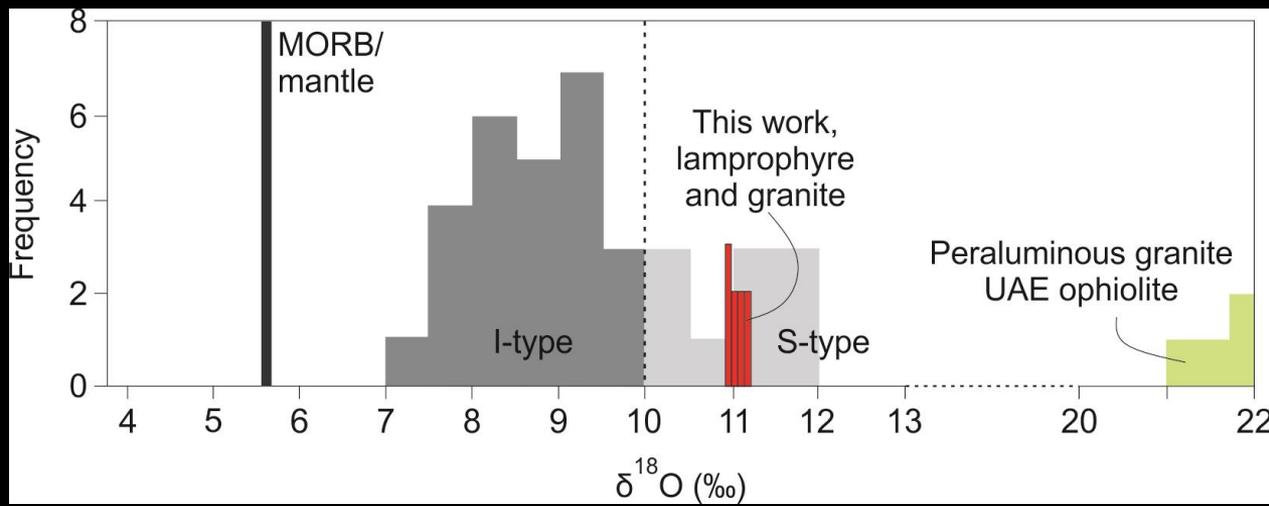
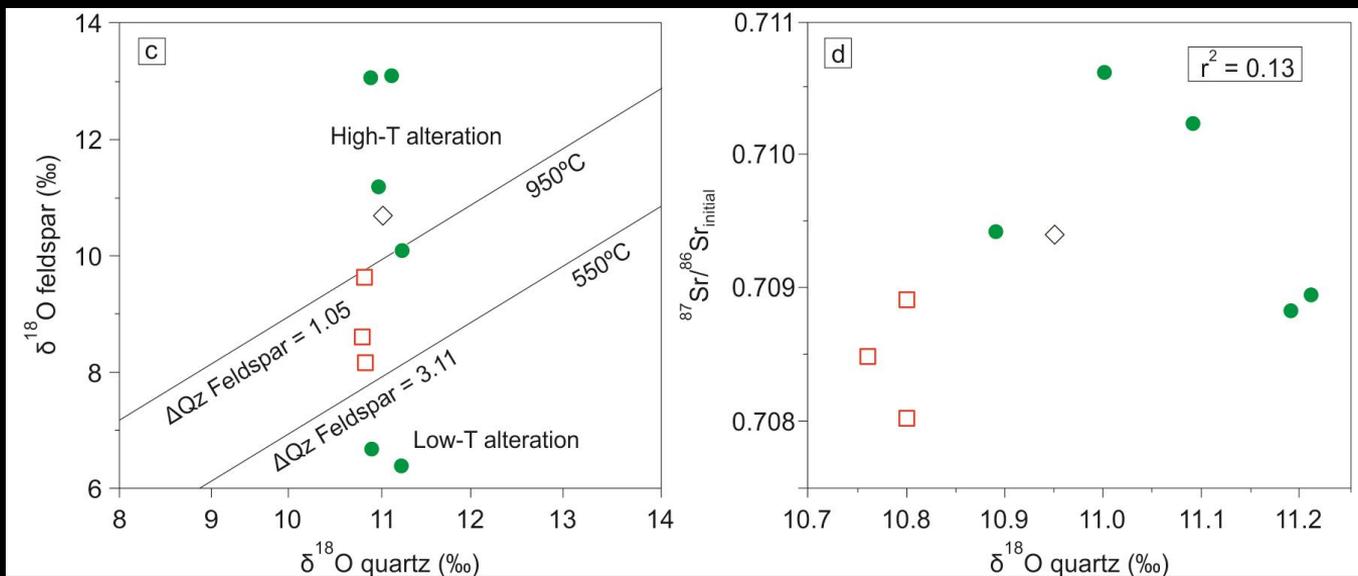
- ✓ Subchondritic  $\epsilon_{\text{Hf}}$  values;
- ✓ Narrow variation with more than 90% of the data plotting between -5 and -1.6



## ➤ Oxygen isotopes in quartz and feldspar

- Quartz  $\delta^{18}\text{O}$  values vary from 10.76 to 11.21‰ ( $\pm 0.15$ ‰)





# Origin of lamprophyres

## “Mantle like” composition

- Moderate to high compatible element contents:
  - Cr (74 – 572 ppm)
  - Ni (23 – 148 ppm)
  - High Mg# (up to ~ 60)



Mantle-derived source for this component of the composite system

## “Crustal like” composition

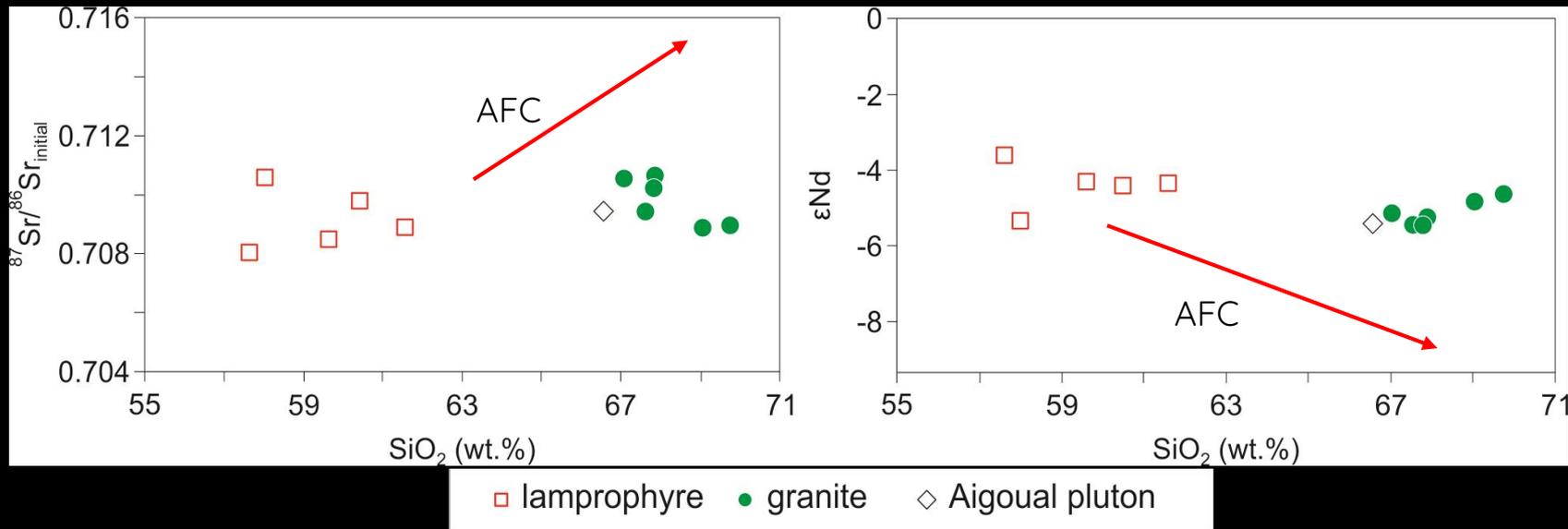
- LILE ( $K_2O$ , Sr, Rb, Ba, Pb, V), LREE;
- Depletion in HFSE such Nb, Ta and Ti in relation to the primitive mantle;
- Subchondritic  $\epsilon Nd$  and  $\epsilon Hf$  values
- $\delta^{18}O$  values that are 5‰ higher than depleted mantle

- The crustal-like signatures of lamprophyres can result from two process:

Assimilation of crust and/or fractional crystallization from a primitive basaltic liquid (AFC process)

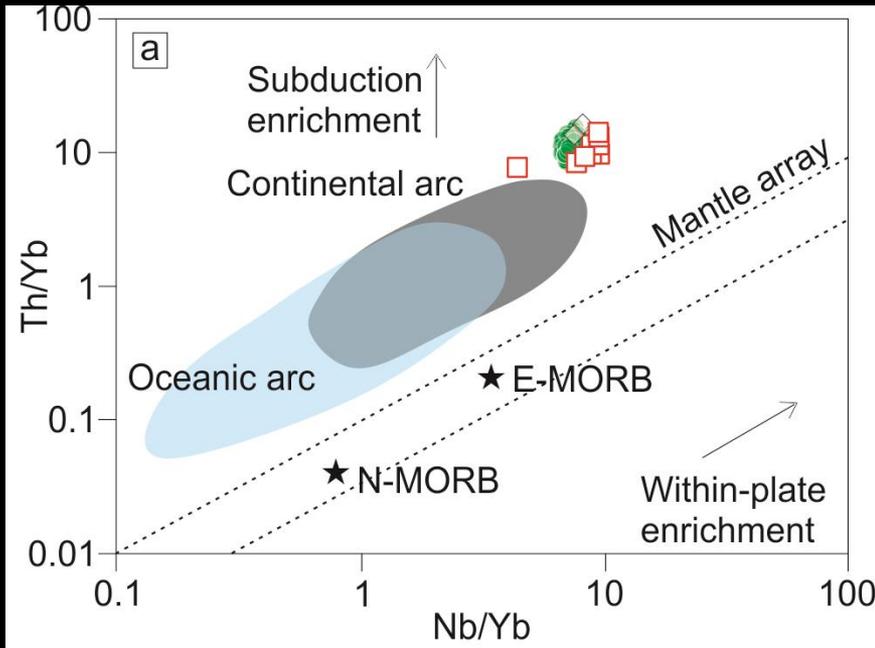
Partial melting of a mantle-source already enriched and metasomatized by fluids or magma derived from crustal material during the subduction stage preceding continental collision

# Assimilation of crust and/or fractional crystallization from a primitive basaltic liquid (AFC process)

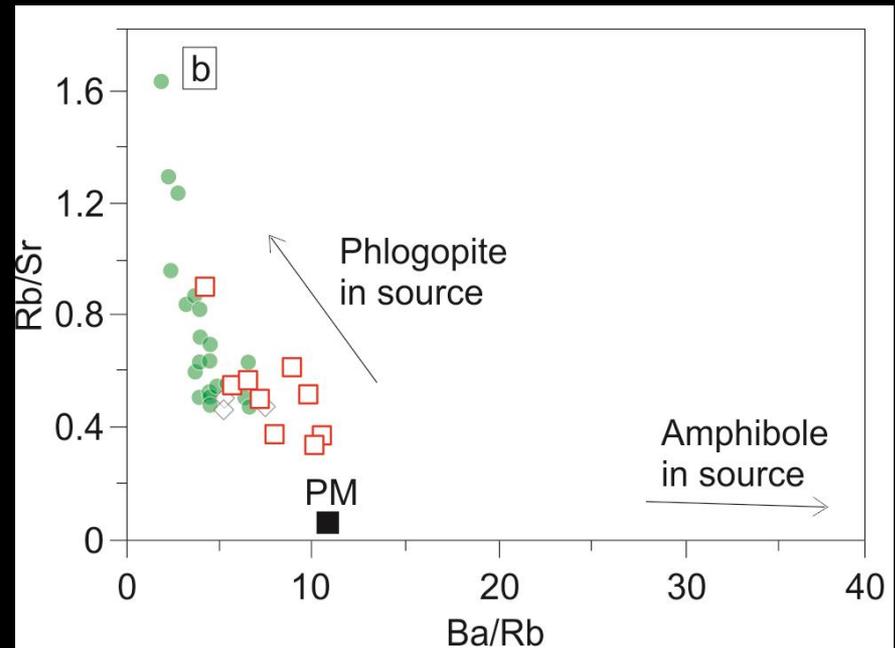


- No correlation between Mg#, Cr, Ba, Ni and  $^{87}\text{Sr}/^{86}\text{Sr}_i$ ,  $\epsilon\text{Nd}$ ,  $\epsilon\text{Hf}$  - isotopic enrichment did not result from crustal contamination;
- Lamprophyres and others high K-Mg are ubiquitous in FMC and intrude several lithologies from the nappe pile. All these rocks display similar signatures to lamprophyres reported in this work and point to **enriched mantle as the most likely source**

Partial melting of a **mantle-source already enriched and metasomatized** by fluids or magma derived from crustal material during the subduction stage preceding continental collision



Magma derived from mantle sources **modified by subducted components** usually differ from MORBs and OIB by enrichment in Th and Nb compared to Yb



Higher Rb/Sr and lower Ba/Rb ratios strongly indicate that **phlogopite** was the dominant **hydrous phase in the source**

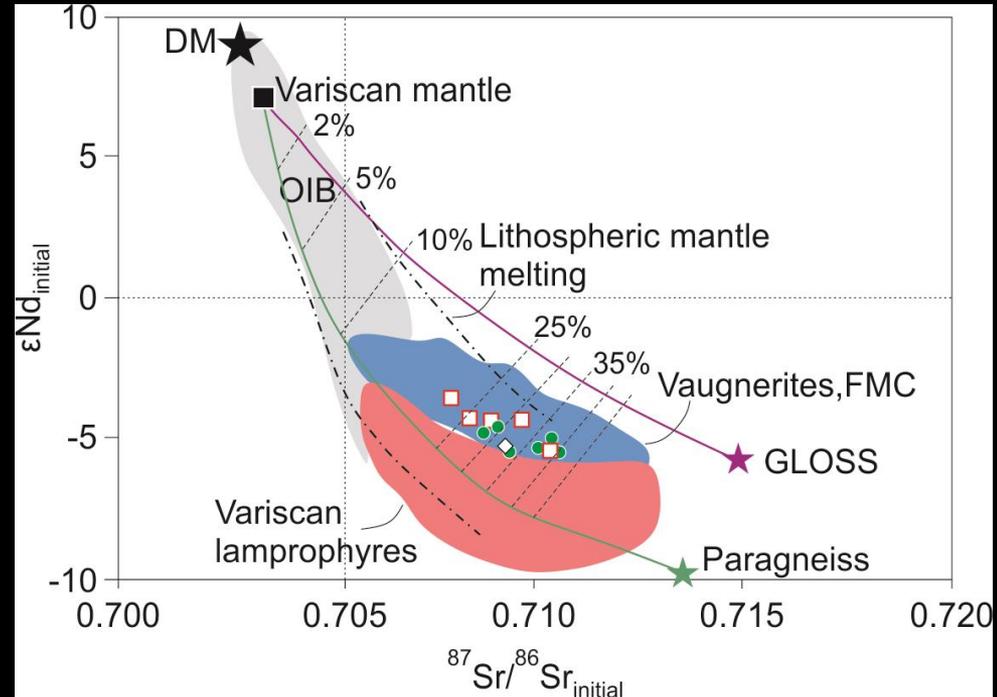
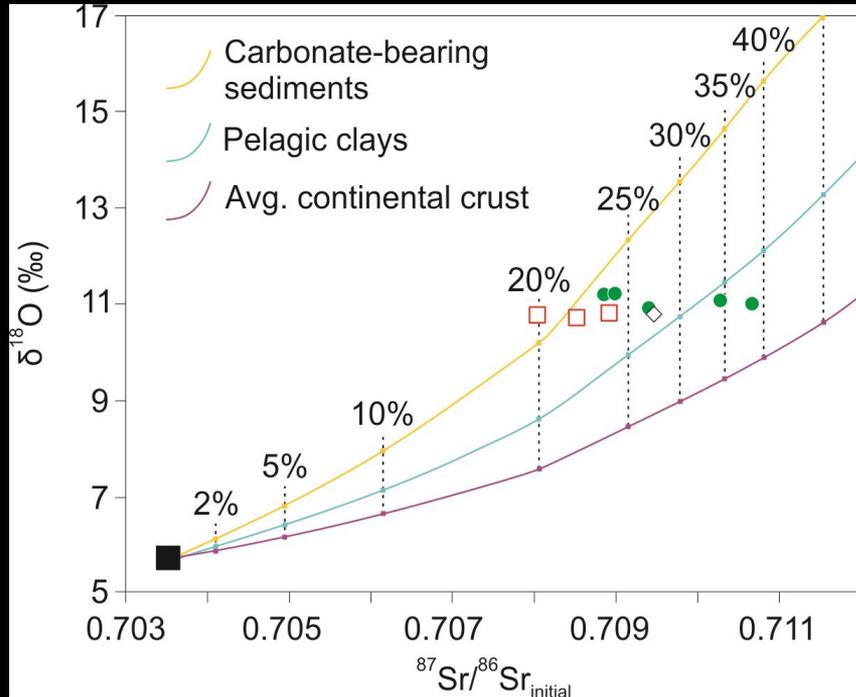
- ✓ Enrichment in Th and Nb compared to Yb;
- ✓ Phlogopite was the dominant hydrous phase in the source;
- ✓ Similar to modern sediments (Global subducted sediments – GLOSS; Plank and Langmuir 1998) - jagged spidergrams marked by low Nb-Ta-Ti concentrations and elevated Th-U and Pb;
- ✓ Low  $^{87}\text{Sr}/^{86}\text{Sr}_i$  ratios are consistent with mantle enrichment during the Silurian-Devonian subduction – Variscan belt



Enriched lithospheric mantle source

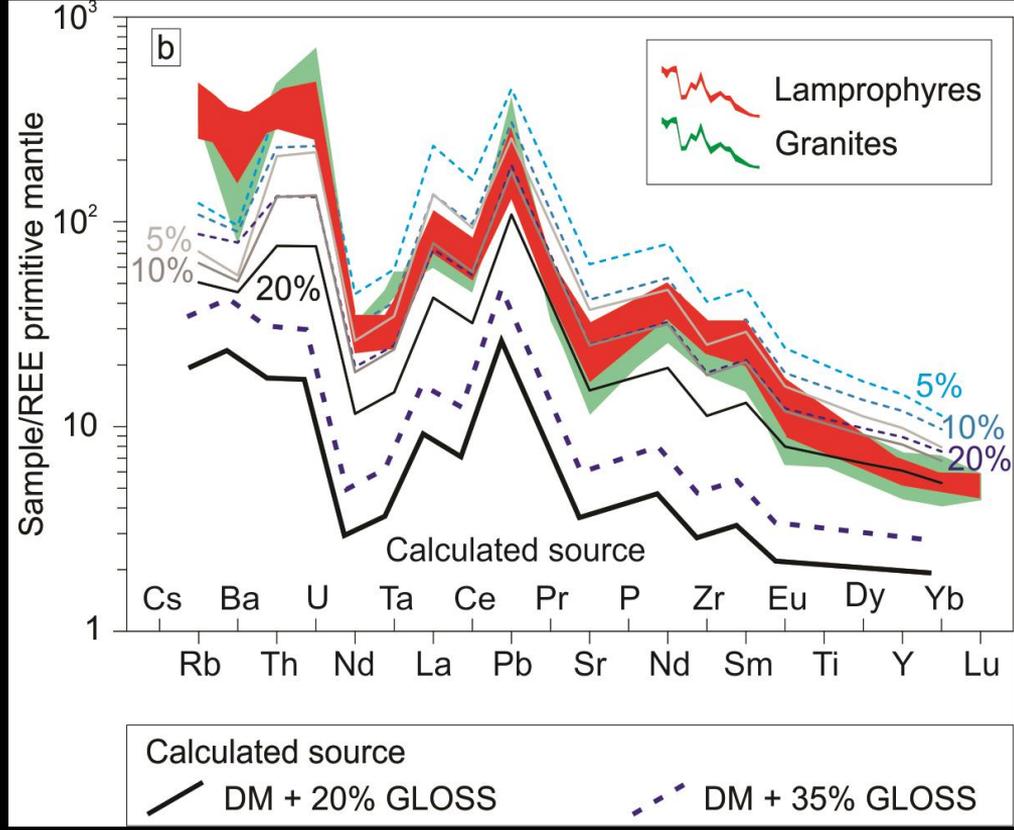
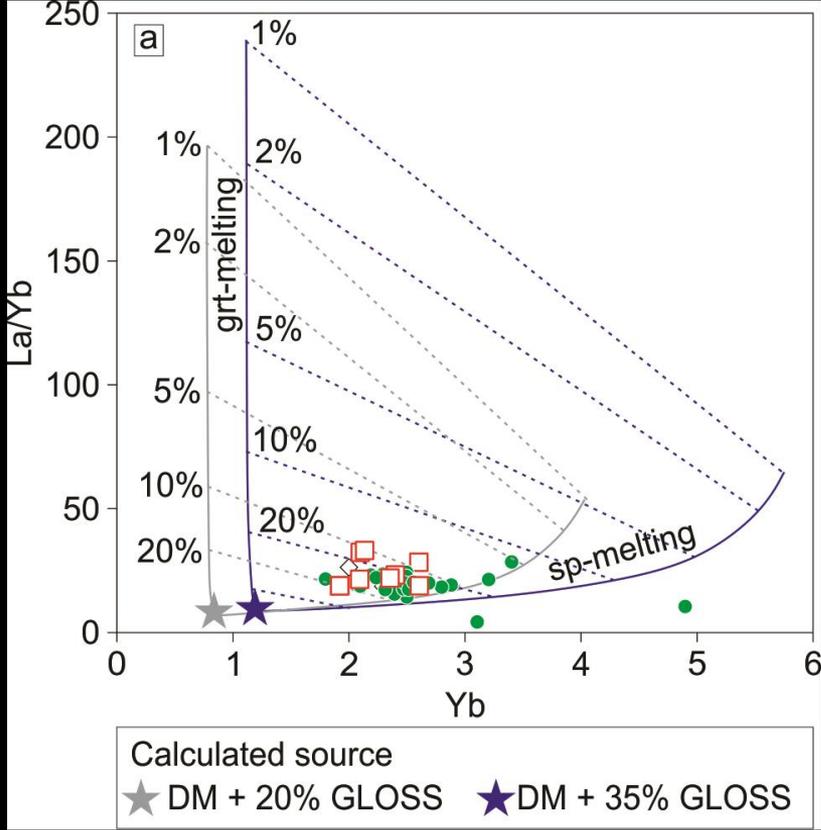
# Nature of the mantle source?

- Modelling using  $^{87}\text{Sr}/^{86}\text{Sr}_{\text{initial}}$  values,  $\epsilon\text{Nd}$  and oxygen isotopes in quartz



□ lamprophyre   
 ● granite   
 ◇ Aigoual pluton

# Partial melting conditions



Shallow depth in spinel-lherzolite facies  
source in the absence of garnet

# Origin of granites

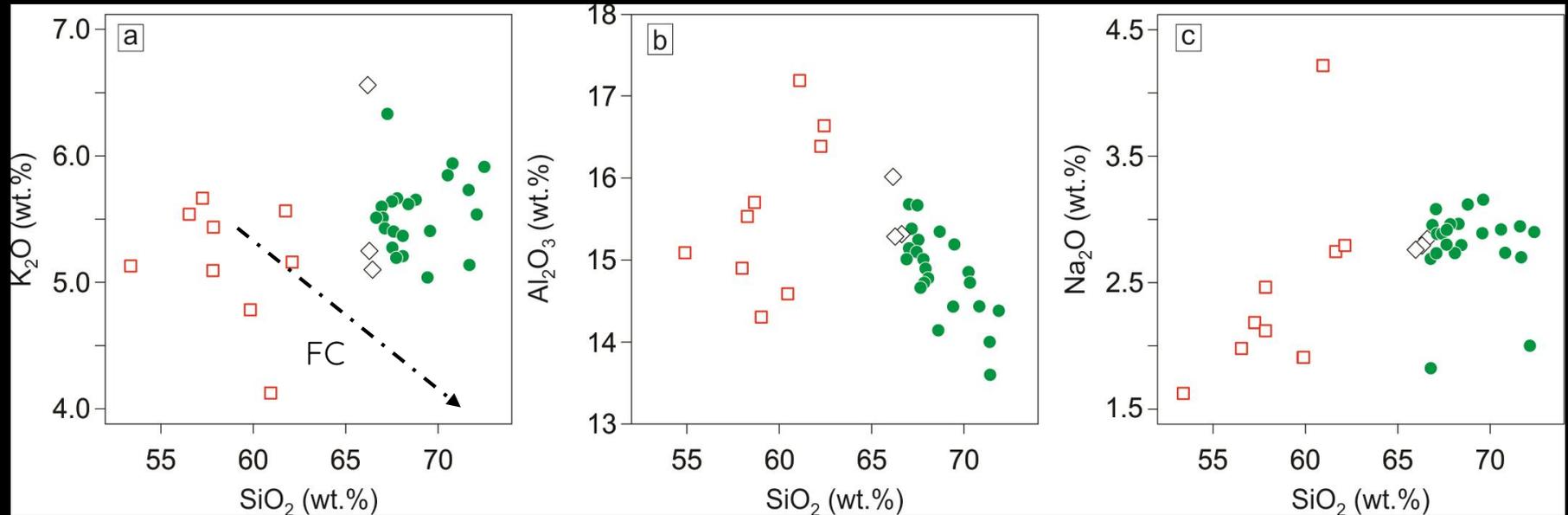
- ✓ The studied dykes are composite and include co-magmatic lamprophyric and granite components;
- ✓ Granites and lamprophyres show a mixture of “crust-like” and “mantle-like” geochemical features;
- ✓ Similar isotopic signatures.

1. The lamprophyres and the granites are related by differentiation processes such as **fractional crystallization**, although remelting of underplated lamprophyres would result in broadly the same patterns; **the lamprophyres differentiated during intrusion into the crust**;

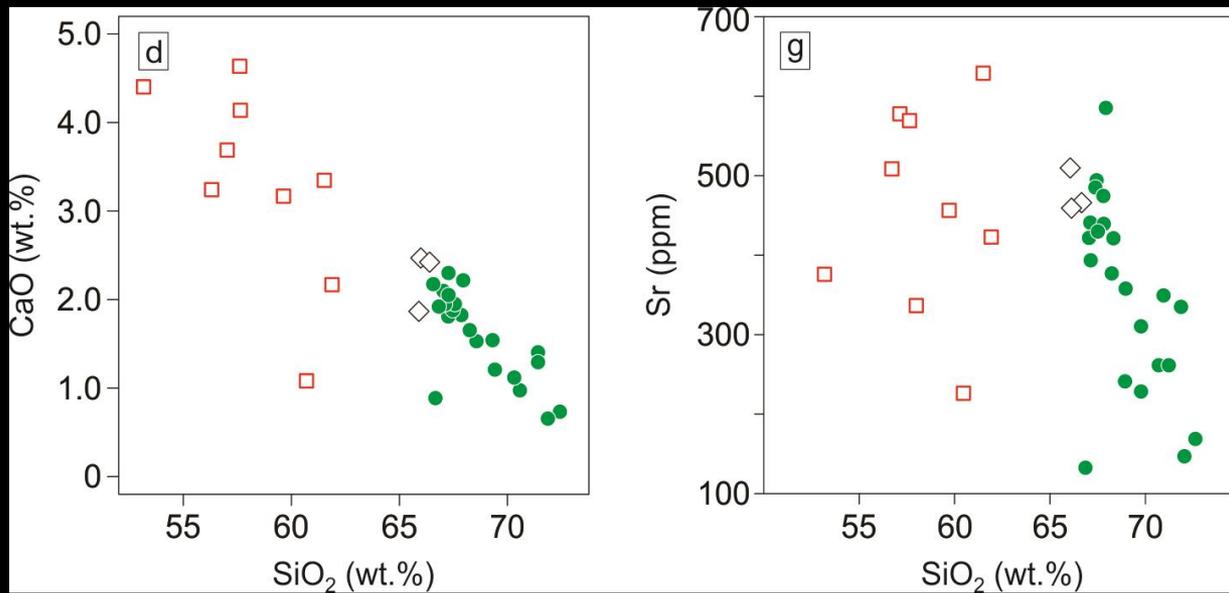
2. The lamprophyres and the granites represent **distinct magmas**, an enriched-mantle derived magma for the lamprophyres and a crust-derived magma for the granites, that **mingle and mix at the emplacement site**

3. The granites and the lamprophyres are both direct **products of melting of the enriched mantle source** but in different proportions

# 1. The lamprophyres and the granites are related by differentiation processes such as fractional crystallization (FC)



The increase in volume and size of megacrysts of K-feldspar in the granitic component demonstrate that a considerable decrease in K<sub>2</sub>O is expected in fractional crystallization in high-K magmas.



- ✓ No cumulate textures are observed in lamprophyres;
- ✓ The high Cr (30 – 140 ppm; 80.9 on average and in some cases over 100 ppm in a granite with SiO<sub>2</sub> = 70 wt.%) and Ni (14 – 40 ppm; 24.4 on average) content in the granites are inconsistent with fractional crystallization playing an important role in the formation of granite.

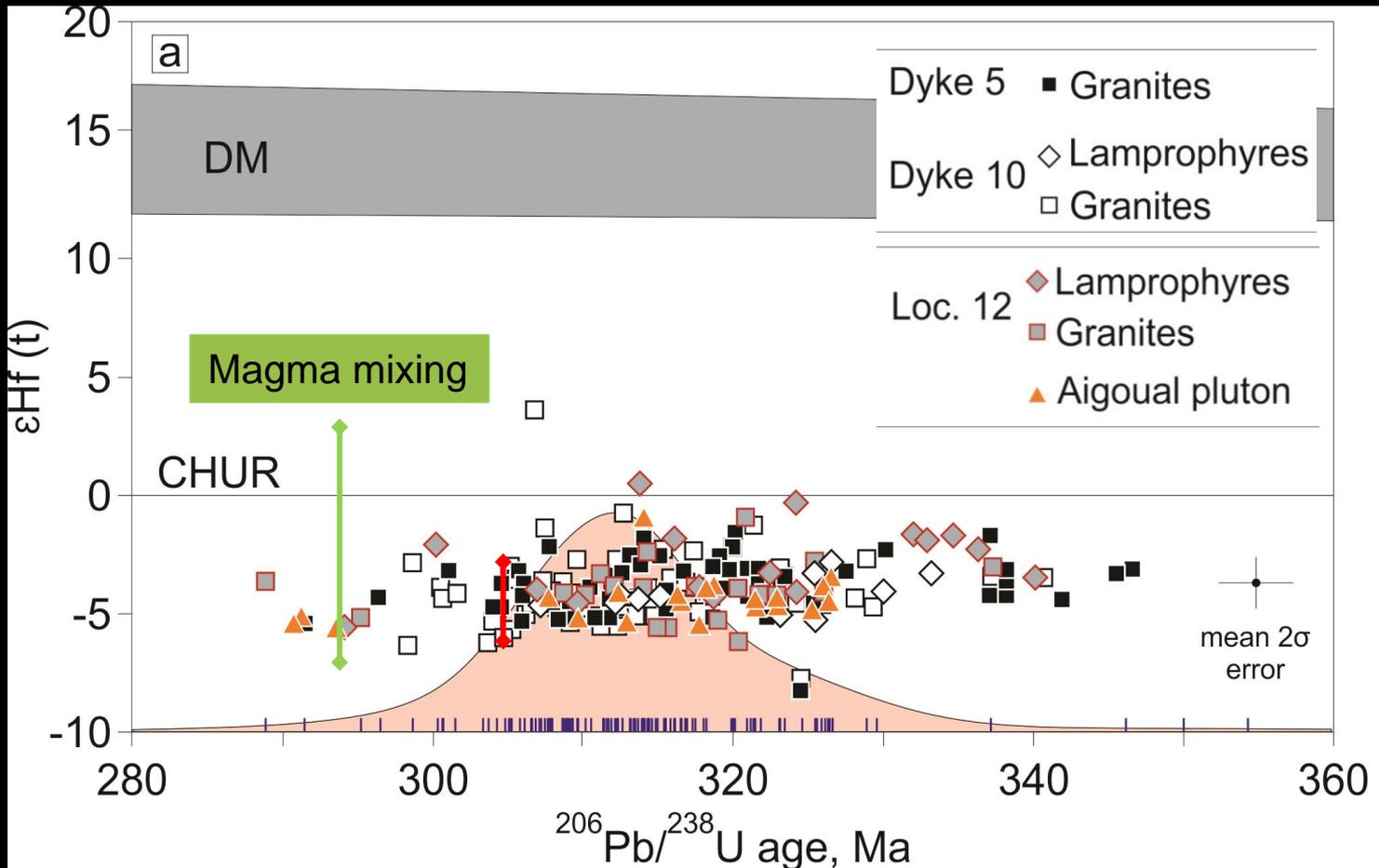
1. The lamprophyres and the granites are related by differentiation processes such as **fractional crystallization**

2. The lamprophyres and the granites represent **distinct magmas**, an enriched-mantle derived magma for the lamprophyres and a crust-derived magma for the granites, that **mingle and mix at the emplacement site**

- ✓ Petrological evidence - ca. 10% of zircon xenocrysts in granites from composite dykes;
- ✓ Corroded of quartz xenocrysts mantled by carbonate and amphibole in lamprophyre;
- ✓ Linear trends observed in many (but not all) binary plots between lamprophyre and granite might indicate magma mixing between these two end members, however, similar trends are observed in fractional crystallization or partial melting processes

**Radiogenic isotopes** are generally not fractionated during melting or crystallization and are **the best tools to decipher open-system processes** such as magma mixing

# ➤ In-situ Lu-Hf isotopic analysis in zircon

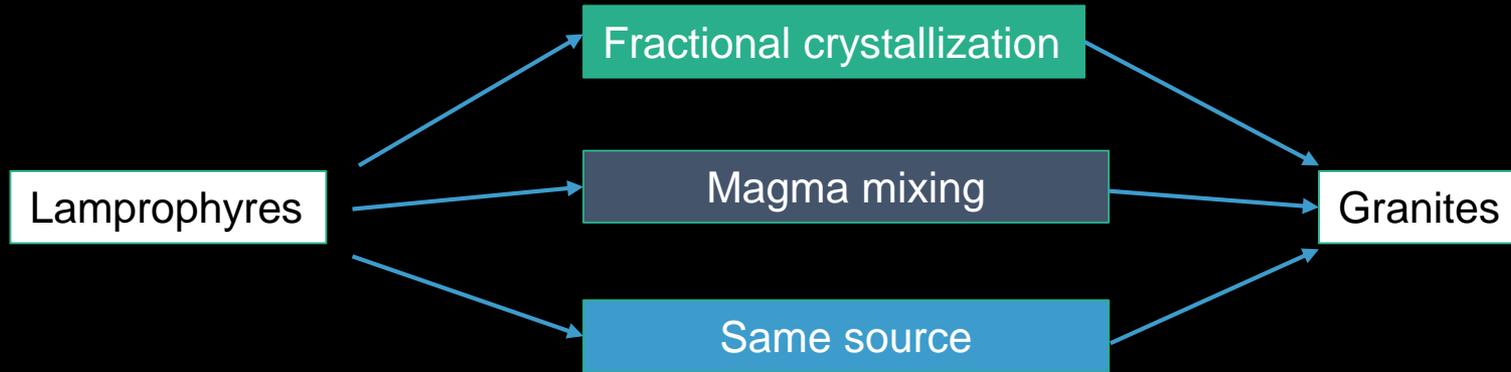


- ✓ The  $\epsilon_{\text{Hf}}$  values from lamprophyres and granites are similar and overlap, pointing that mixing within different and contrasting magmas did not play an important role;
- ✓ The narrow range in  $\delta^{18}\text{O}$  values between lamprophyres and granites (less than 0.5‰) is more consistent with a source related variation or homogenization of the magma than mixing between two different and contrasting magmas;
- ✓ Although crustal-derived melts are regionally widespread in FMC, the Aigoual pluton lacks purely crustal granites, e.g. in the form of peraluminous cordierite- and muscovite-bearing rocks that are ubiquitous in the FMC;
- ✓ The lack of well-defined mixing trends for major elements, trace elements and isotopes alike suggests that magma mixing was not a dominant process;
- ✓ A minimal degree of assimilation is likely.

2. The lamprophyres and the granites represent **distinct magmas**, an enriched-mantle derived magma for the lamprophyres and a crust-derived magma for the granites, that **mingle and mix at the emplacement site**

3. The granites and the lamprophyres are both direct **products of melting of the enriched mantle source** but in different proportions

- ✓ Co-magmatic relation between lamprophyres and granites and very similar isotopic signatures;
- ✓ Experimental studies – peridotite + sediments or peridotite/lherzolite + granitic melt → post-collisional K-rich magmatism;
- ✓ Recent studies (Förster et al. 2020, 2021) → mixing between peridotite and sediment can reach high SiO<sub>2</sub> contents (> 67%) and major and trace element compositions matching those in this work.

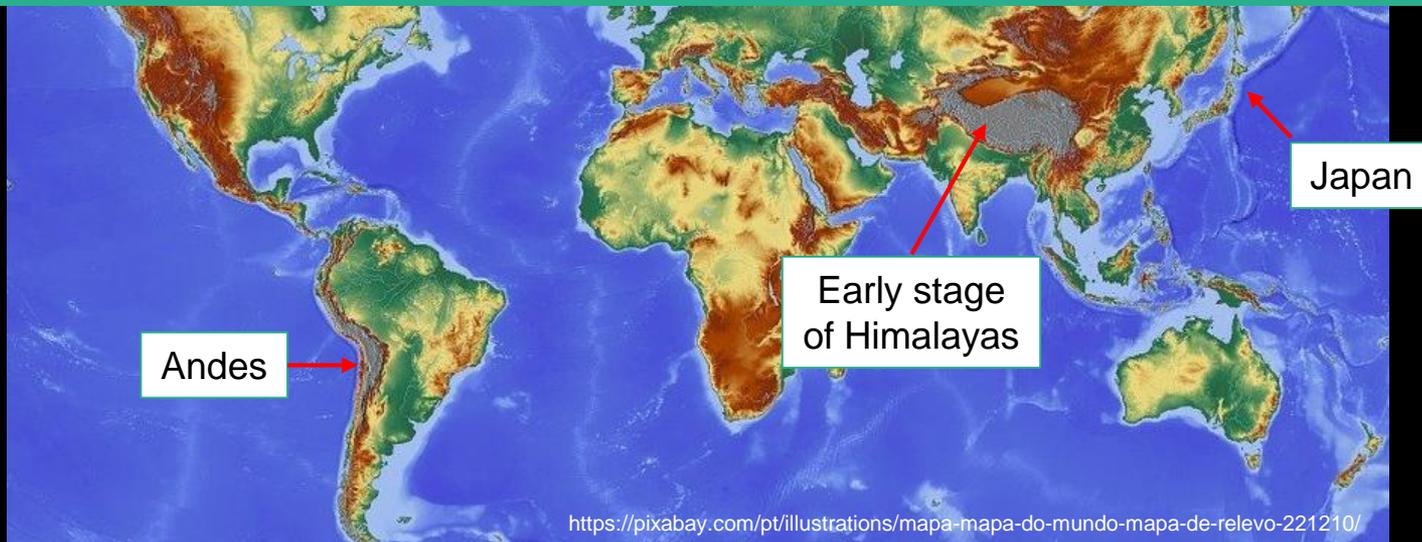


- ✓ Both lamprophyres and granites display evidence of **crustal recycling** in different ways and **addition of new mantle-derived** material as seen in the high contents of **compatible elements such Cr, Ni, MgO and FeO**.

# Implications for crustal growth

- The crust grows when new mantle-derived material is added to the continent

Arc settings - large amounts of mantle-derived mafic and intermediate magmas (and their differentiates) are added to the crust



However, this does not necessarily imply significant **long-term crustal-growth**, because **arc settings display poor preservation potential**. A high proportion of the generated crust may be recycled back to the mantle shortly after formation (Condie 2014; Korenaga 2018; Scholl and von Huene 2009; Stern 2011)

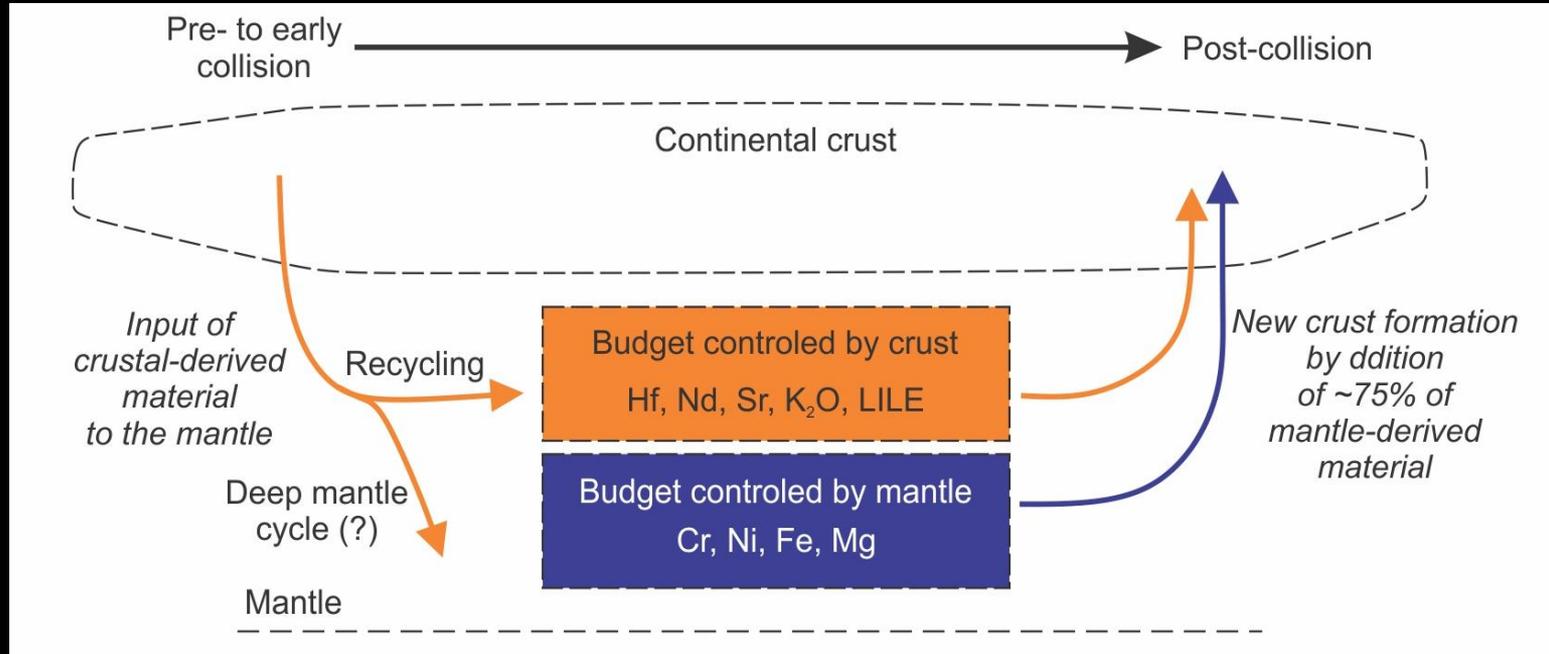
## Post-collisional setting

- ✓ Less efficient than arc settings in producing new continental crust;
- ✓ Highest potential for preservation in the geological record (Hawkesworth et al. 2009, 2010; Spencer et al. 2015).

Post-collisional settings may represent significant contributions for the long-term crustal growth

## Post-collisional setting

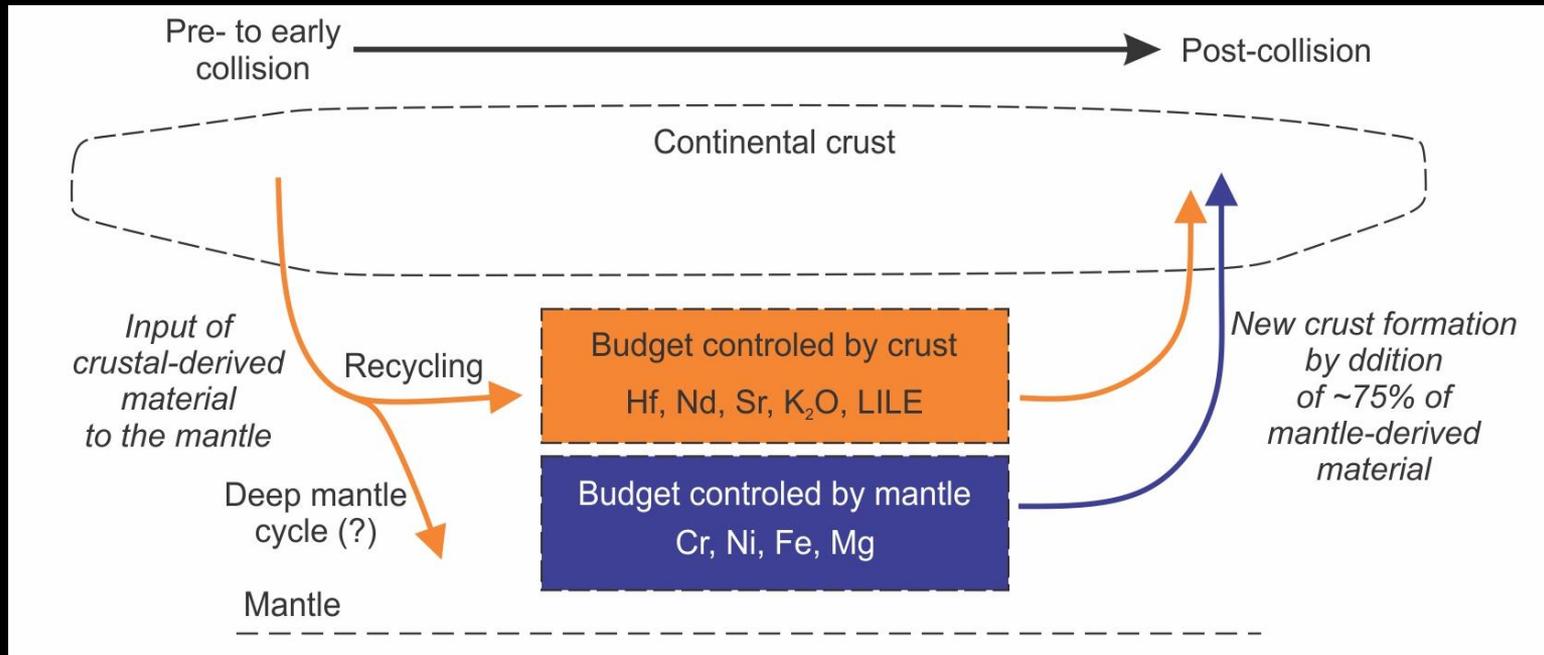
- ✓ Addition of new mantle-derived material is not identified by usual isotopic systems such as Rb-Sr, Sm-Nd and Lu-Hf



~25% of crustal material is recycled back to the continental crust

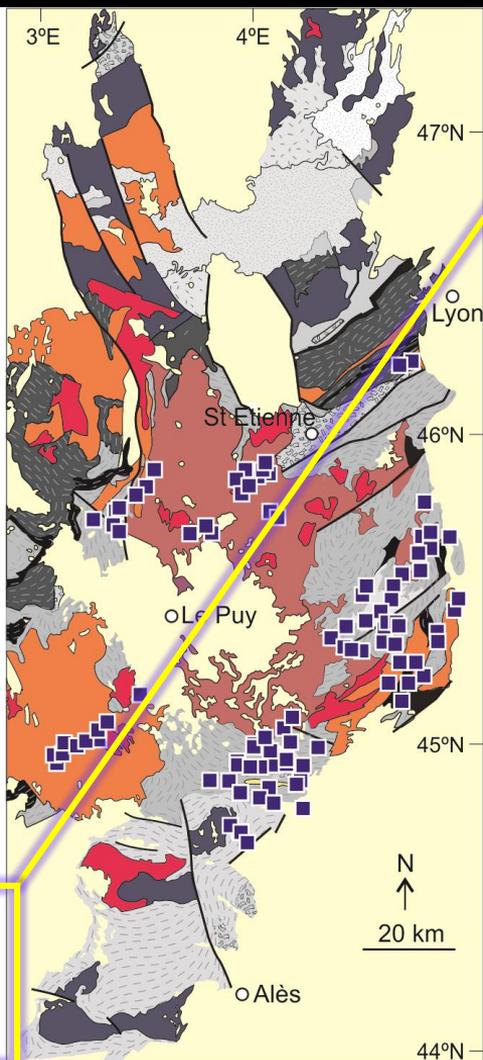
~75% of mantle-derived material is added to the crust corresponding new crust formation

The mantle-derived material is evidenced by the mafic character and the high contents of compatible elements that ultimately came from the mantle.



### Legend

- Major fault / shear zone
- Post-Permian cover
- Variscan units and nappe stack
  - Upper Carboniferous and Permian detrital basin
  - Lower Carboniferous volcano-sedimentary rock (a): associated granophyre
  - Devonian volcano-sedimentary rock (Brévenne unit)
  - Upper Gneiss Unit (a): Leptyno-amphibolitic complex
  - Lower Gneiss Unit
  - Velay migmatite, granite
  - Parautochthonous (Cévennes schists)
- Variscan plutonic rocks
  - Bt ± Ms peraluminous leucogranite (MPG)
  - Bt ± Crd peraluminous granite and granodiorite (CPG)
  - Bt ± Am peraluminous to metaluminous granite to quartz diorite (KCG)
  - High K-Mg rocks «Vaugnerites»



High K-Mg rocks + KCG plutons

35% of the outcrop surface of granites in FMC

+

Geochemical modelling display between 65 – 80% of the material came from the mantle



22 – 28% of the magmatic material in FMC came from the mantle



Crustal growth

# Summary

- Introduction
- Objectives and Methodology
- Results and Discussion

Chapter II – paper in preparation to submission

Chapter III – general discussion

- **Conclusions**

# Conclusions

- 1) Lamprophyres and granites are co-magmatic and coeval with the emplacement of Aigoual pluton – between 311 and 313 Ma;
- 2) Lamprophyres display high contents of compatible (Cr, Ni, Fe, Mg) and incompatible elements (LILE,  $K_2O$ , Ba, Pb, Sr) and crustal-like signatures in both radiogenic and stable isotopic systems;

Partial melting and mixing modelling suggest this dual geochemical signature results from 10 – 20% of partial melting in the spinel-lherzolite stability field of a source composed by mixing between 20 – 35% of sediments and 65 – 80% peridotite;

- 3) Granites display trace elements and isotopic signatures alike lamprophyres and have a mantle-derived component involved in their petrogenesis. The intimate relation (physically and geochemically) between lamprophyre and granite in a system of composite dykes of northern Aigoual pluton provide constraints about the mantle contribution in the granites from FMC;
- 4) Although the isotopic signatures being broadly controlled by the crustal component in the mantle, approximately 70% of the material comes from the mantle and thus corresponds new crust formation;
- 5) Lamprophyres and granites from composite dykes represent significant addition of new crust in a post-collisional setting.

# References

- Condie KC (2014). Growth of continental crust: a balance between preservation and recycling. *Mineralogical Magazine* 78(3):623-637.
- Couzinié S, Laurent O, Moyen JF, Zeh A, Bouilhol P, Villaros A (2016). Post-collisional magmatism: Crustal growth not identified by zircon Hf-O isotopes. *Earth and Planetary Science Letters* 456:182-195.
- Couzinié S (2018). Evolution of the continental crust and significance of the zircon record, a case study from the French Massif Central. PhD thesis, Stellenbosch University and Saint-Etienne University. 436 pp.
- Hawkesworth C, Dhuime B, Pietranik AB, Cawood PA, Kemp AIS, Storey CD (2010). The generation and evolution of the continental crust. *Journal of the Geological Society*, 167(2):229-248.
- Hawkesworth C, Cawood P, Kemp T, Storey C, Dhuime, B (2009). Geochemistry: A matter of preservation. *Science*, 323:49-50.
- Korenaga J (2018). Crustal evolution and mantle dynamics through Earth history. *Philosophical Transactions Royal Society A* 376:20170408. <http://dx.doi.org/10.1098/rsta.2017.0408>.
- Laurent O, Couzinié S, Zeh A, Vanderhaeghe O, Moyen JF, Villaros A, Gardien V, Chelle-Michou C (2017). Protracted, coeval crust and mantle melting during Variscan late-orogenic evolution: U–Pb dating in the eastern French Massif Central. *Int J Earth Sci (Geol Rundsch)* 106:421–451.
- McDonough WF and Sun SS (1995). The composition of Earth. *Chemical Geology* 120:223-253.
- Moyen JF, Laurent O, Chelle-Michou C, Couzinie S, Vanderhaeghe O, Zeh. A, Villaros S, Gardien V (2017). Collision vs. subduction-related magmatism: Two contrasting ways of granite formation and implications for crustal growth. *Lithos* 277:154-177.

Plank T and Langmuir CH (1998). The chemical composition of subducting sediment and its consequences for the crust and mantle. *Chemical Geology* 145:325-394.

Scholl DW and von Huene R (2009). Implications of estimated magmatic additions and recycling losses at the subduction zones of accretionary (non-collisional) and collisional (suturing) orogens. Geological Society, London, Special Publications 318(1):105-125.

Shand SJ (1943). *Eruptive Rocks. Their Genesis, Composition, Classification, and Their Relation to Ore-Deposits with a Chapter on Meteorite.* John Wiley & Sons, New-York.

Spencer CJ, Cawood PA, Hawkesworth CJ, Prave AR, Roberts NMW, Horstwood MSA, Whitehouse MJ, EIMF (2015). Generation and preservation of continental crust in the Grenville Orogeny. *Geoscience Frontiers* 6 (3):357-372.

Stern CR (2011). Subduction erosion: rates, mechanisms, and its role in arc magmatism and the evolution of the continental crust and mantle. *Gondwana Research* 20(2):284-308.

Su HM, Jiang SY, Zhang DY, Wu XK (2017). Partial Melting of Subducted Sediments Produced Early Mesozoic Calc-alkaline Lamprophyres from Northern Guangxi

Talbot JY, Chen Y, Faure M (2005). A magnetic fabric study of the Aigoual–Saint Guiral–Liron granite pluton (French Massif Central) and relationships with its associated dikes. *Journal of Geophysical Research* 110:B121106.

von Raumer JF, Finger F, Vesela P, Stampfli GM (2014). Durbanchites-Vaugnerites – a geodynamic marker in the central European Variscan orogen. *Terra Nova* 26:85-95.

Thank you!



Obrigada!

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Mantle and crust interaction in post-collisional setting,  
a case study of lamprophyric-granitic composite dykes  
of northern Aigoual pluton, French Massif Central

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